

Article

An Uncertainty-Aware Assessment Framework for Multi-Glider Coordinated Observation of Mesoscale Eddies

Dongyang Xue ¹, Fang Liu ², Yaqiang Zhu ^{1,*}, Xuehao Wang ¹, Shuai Li ¹, Shufeng Li ³ and Peng Wang ⁴

¹ School of Mechanical Engineering, Tianjin University of Commerce, Tianjin 300134, China; dongyangxue@tjcu.edu.cn (D.X.); xuehaowang@tjcu.edu.cn (X.W.); lishuai1@tjcu.edu.cn (S.L.)

² Yichang Research Institute of Testing Technology, Yichang 443003, China; liufanglfaq163@163.com

³ School of Aeronautical Engineering, Civil Aviation University of China, Tianjin 300300, China; sf_li@cauc.edu.cn

⁴ School of Mechanical Engineering, Tianjin University, Tianjin 300354, China; peng_wang@tju.edu.cn

* Correspondence: zyq@tjcu.edu.cn

Abstract

Autonomous underwater glider fleets are increasingly deployed to observe mesoscale eddies, yet a methodology for evaluating observation quality under uncertainty remains lacking. This paper presents SCOPE, a framework integrating uncertainty propagation, multi-dimensional assessment, and objective metric selection. A 27-metric evaluation system spanning seven quality categories is constructed; an Adaptive Core Metric Selection (ACMS) algorithm compresses these to a compact core subset. Two-stage sensitivity analysis identifies the velocity ratio as the dominant parameter with a non-monotonic effect. The optimal velocity ratio decreases with eddy intensity in both an analytical model (0.58 to 0.47) and four real HYCOM eddies (2.35 to 0.35). ACMS converges on all four real eddies, and parameter rankings are consistent across environments. An ablation experiment in the analytical model validates the effectiveness of the two sensitivity-derived design rules (25% and 45.8%) gap closure and reveals their non-additive interaction. SCOPE offers a framework and diagnostic workflow for uncertainty-aware evaluation of autonomous ocean observation systems.

Keywords: mesoscale eddy; underwater glider; coordinated observation; uncertainty quantification; sensitivity analysis; performance evaluation; SCOPE framework

1. Introduction

Mesoscale eddies are ubiquitous rotating structures in the global ocean, typically spanning 50–300 km in diameter and persisting for weeks to months [1,2]. These coherent vortices play a critical role in the redistribution of oceanic energy, heat, salt, and biogeochemical tracers, significantly influencing large-scale ocean circulation, regional climate variability, and marine ecosystem dynamics [3–6]. Understanding their three-dimensional structure is essential for improving ocean prediction models and quantifying their contribution to global energy budgets [2,7]. Satellite remote sensing detects surface signatures but lacks subsurface information [1,8]. Subsurface profiling platforms, including Argo floats, provide depth-resolved measurements but lack the spatial resolution and maneuverability to resolve the fine-scale internal structure of individual eddies [9]. This observational gap motivates the deployment of coordinated mobile platforms capable of three-dimensional sampling within targeted eddy features.



Academic Editor: Angelo Rubino

Received: 26 May 2026

Revised: 28 June 2026

Accepted: 29 June 2026

Published: 30 June 2026

Copyright: © 2026 by the authors.

Licensee MDPI, Basel, Switzerland.

This article is an open access article distributed under the terms and conditions of the [Creative Commons Attribution \(CC BY\)](https://creativecommons.org/licenses/by/4.0/) license.

Autonomous underwater gliders have emerged as a promising platform for in situ eddy observation, owing to their extended endurance, depth-profiling capability, and programmable navigation [10–12]. Recent studies have demonstrated their effectiveness in capturing eddy internal structures [13,14]. Deploying multiple gliders in a coordinated network has been validated in field experiments [15] and enables systematic sampling of an eddy's spatial structure through cross-sectional transects [16,17]. Several path planning strategies have been proposed, including butterfly patterns that traverse the eddy core along intersecting diameters [18], zigzag lawn-mower patterns for area coverage [19] and adaptive approaches such as heading control for eddy tracking [20] and real-time waypoint adjustment [21,22]. These methods optimize controllable parameters, such as glider speed, heading angle, centerline orientation, and intersection geometry under the assumption of a known, deterministic eddy environment [23]. However, real ocean eddies are inherently variable: their intensity, vertical structure, and ambient current conditions cannot be precisely known prior to deployment. Whether parameters optimized for one eddy condition—glider speed, in particular—remain effective across a range of eddy intensities is an open question.

Existing studies on glider-based eddy observation typically assume a known environment when evaluating path designs [18,24]. The impact of ocean uncertainty on glider performance has been investigated at the individual platform level. Ocean currents alter glider motion dynamics and trajectory accuracy [25,26], time-varying currents challenge path planning [27,28], and stochastic conditions affect energy consumption [29,30]. These studies address trajectory-level errors under specific uncertainty sources. However, mesoscale eddy observation is a mission-level task, whose outcome is measured by the quality of reconstructed oceanographic fields, spatial coverage, shape and center accuracies. How multi-source uncertainties propagate from individual glider trajectories to these mission-level metrics remains largely unexplored. Systematic frameworks integrating adaptive sampling and uncertainty quantification have been developed for ocean observation [31], but their application to multi-glider eddy observation under multi-source uncertainty has not been addressed. A systematic framework that integrates uncertainty propagation, multi-dimensional evaluation, and parameter sensitivity analysis for glider network observation is lacking.

To address these limitations, this paper proposes the SCOPE (Stochastic Coordinated Observation Performance Evaluation) framework. The principal contributions are:

1. SCOPE framework. A closed-loop simulation framework couples a parameterized multi-scale eddy environment with a kinematic glider model, propagating multi-source uncertainties to a 27-metric evaluation system spanning seven quality categories. An Adaptive Core Metric Selection (ACMS) algorithm compresses the metric set while preserving category coverage.
2. Sensitivity analysis and ablation diagnosis. Correlation and interaction analyses of the 10 uncertain parameters yield a ranked importance hierarchy and identify three physically coupled parameter pairs. Derived variables constructed from these couplings reveal a non-monotonic, intensity-dependent effect of the dominant parameter. Structured ablation experiments validate these findings by decomposing marginal and combined effects of the identified design rules.
3. Methodology transfer to real ocean eddies. The ACMS algorithm and sensitivity ranking methodology are validated on four HYCOM eddies of different intensities. The core metrics and parameter rankings obtained by the same SCOPE methodology are compared between the analytical model and real eddies. The velocity ratio trend observed in the analytical model is reproduced in real eddies, with a systematic shift in magnitude attributable to differences in environmental complexity.

The remainder of this paper is organized as follows. Section 2 presents the SCOPE framework. Section 3 reports the two-stage sensitivity analysis, dose–response characterization, and ablation validation. Section 4 examines the framework’s transferability to real ocean eddies. Section 5 discusses implications and concludes.

2. The SCOPE Framework

2.1. Framework Overview

The SCOPE (Stochastic Coordinated Observation Performance Evaluation) framework is proposed to quantify and diagnose multi-glider observation performance under multi-source uncertainty. Whereas prior work addresses single-glider trajectory errors under individual disturbance sources [25,26], SCOPE integrates environmental, instrumental, and planning uncertainties into a unified propagation pipeline and evaluates their combined effect through a multi-dimensional mission-level metric system. Figure 1 shows its workflow.

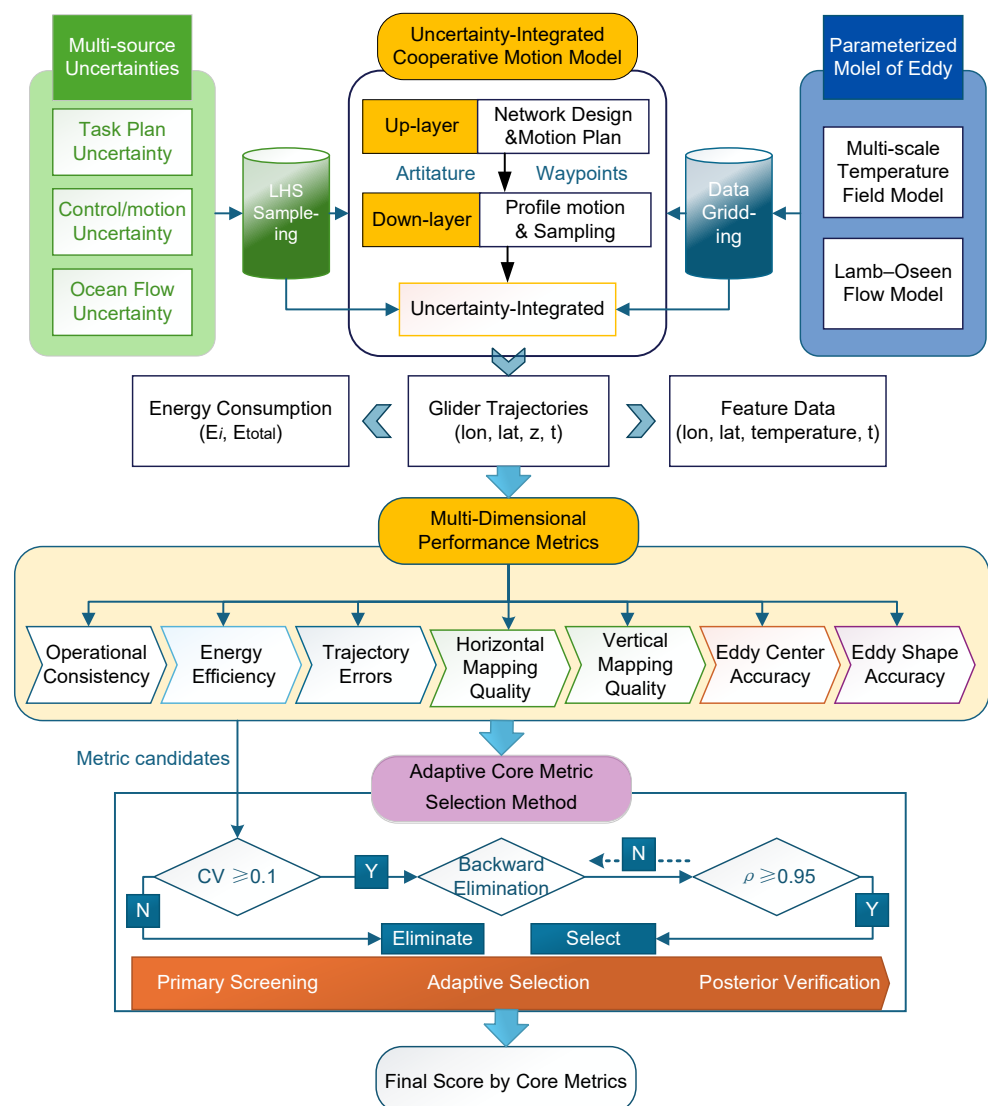


Figure 1. Architecture of the SCOPE framework. Four modules operate sequentially: environment generation, motion simulation, multi-dimensional assessment, and metric selection.

First, a parameterized eddy model (Section 2.2) provides gridded temperature and velocity fields under main uncertain parameters. Second, the Uncertainty-Integrated Coordinated Motion Model (Section 2.3) propagates these uncertainties together with controllable mission parameters and uncontrollable motion factors, outputting glider trajectories and temperature measurements via discrete-time integration. A Monte Carlo experiment is then conducted via Latin Hypercube Sampling (LHS) of the uncertain parameters. Each LHS-sampled mission is scored by 27 metrics across seven quality categories, and the ACMS algorithm compresses them, preserving ranking fidelity (Section 2.4). As shown in the composite score, the framework supports global sensitivity analysis, parameter importance ranking, and characterization of key parameters (Section 3).

2.2. Parameterized Eddy Field

The eddy environment field is constructed by a parameterized velocity field and temperature structure to enable controlled uncertainty propagation and analysis.

2.2.1. Temperature Field

A gridded three-dimensional temperature field (Figure 2a) serves as the observation ground truth. It features a subsurface-intensified warm-core structure consistent with South China Sea anticyclonic eddies, constructed from three superimposed components [32]: a hyperbolic-tangent thermocline background, a warm-core anomaly, and sub-mesoscale perturbations. The dominant warm-core anomaly is modeled as

$$T_{\text{eddy}}(r, z) = A_T \cdot e^{-\beta r_n^2} \cdot e^{-(z-z_T)^2/(2H_T^2)} \cdot [1 + \gamma e^{-(r_n-1)^2/(2\sigma_e^2)}] \quad (1)$$

where $r_n = r/r_{\text{edge}}(\phi)$ is the normalized radial coordinate, z_T is the depth of maximum thermal anomaly, H_T is the vertical e-folding half-width, β is the radial decay rate of the temperature anomaly, γ is the amplitude of the thermal front enhancement at the eddy boundary, and σ_e is the width of the front enhancement. The final factor (γ, σ_e) produces enhanced thermal fronts at the eddy boundary [33]. Sub-mesoscale perturbations are added as multi-frequency sinusoidal fluctuations (amplitudes 0.2–0.3 °C) to represent filamentary structures and internal waves [34]. Gaussian observation noise ($\sigma = 0.08$ °C) is superimposed. The eddy boundary is defined as a perturbed ellipse:

$$r_{\text{edge}}(\phi) = r_{\text{ellipse}}(\phi) + A_{w3} \sin(3\phi) + A_{w8} \sin(8\phi + \pi/4) + \epsilon_r(\phi) \quad (2)$$

where $r_{\text{ellipse}}(\phi) = ab / \sqrt{(b \cos \phi)^2 + (a \sin \phi)^2}$ is the standard ellipse equation, with semi-axes $a = 60$ km (aligned with the major axis), $b = 51$ km (minor axis), $A_{w3} = 5$ km, $A_{w8} = 2$ km, and $\epsilon_r \sim \mathcal{N}(0, 1.5^2)$ km.

2.2.2. Flow Field

The Flow field (Figure 2b) uses a Lamb–Oseen vortex model [35]: The tangential velocity at radial distance r from the eddy center is

$$V_\theta(r) = V_{\text{max}} \cdot \frac{r}{R_{\text{max}}} \cdot \exp\left[\frac{1}{2}\left(1 - \frac{r^2}{R_{\text{max}}^2}\right)\right] \quad (3)$$

where V_{max} is the maximum tangential velocity at radius R_{max} . Vertical decay follows $f(z) = \exp[-(z - z_m)^2/(2\sigma_z^2)]$ with depth of maximum swirl z_m and vertical e-folding scale σ_z . A spatially uniform background current U_c with direction θ_c is superimposed to represent mesoscale advection:

$$\mathbf{u}(r, \phi, z) = V_\theta(r) f(z) \hat{\mathbf{e}}_\phi + U_c(\cos \theta_c, \sin \theta_c) \quad (4)$$

Vertical velocity (negligible relative to horizontal components) is neglected to simplify computations.

This parameterized formulation allows the key environmental variables— $V_{\max}$, z_m , U_c , and θ_c —to be treated as controllable inputs for uncertainty propagation. By varying these parameters within their prescribed ranges (Table 1), the model generates an ensemble of eddy realizations that supports Latin Hypercube Sampling and Monte Carlo experiments.

Table 1. Physical parameters of the parameterized environmental field.

Symbol	Description	Value	Unit
$R_{\max}$	Radius of maximum velocity	45	km
σ_z	Vertical e-folding scale	400	m
T_{surf}	Surface temperature	28	°C
T_{deep}	Deep temperature	8	°C
z_{th}	Thermocline center depth	150	m
A_T	Peak thermal anomaly	4.0	°C
z_T	Depth of maximum thermal anomaly	150	m
H_T	Vertical e-folding half-width	200	m
—	Extent	$2^\circ \times 2^\circ$	—
Δx	Horizontal resolution	$1/24^\circ \approx 4.6$	km
Δz	Vertical resolution	20	m
$D_{\max}$	Maximum depth	1000	m

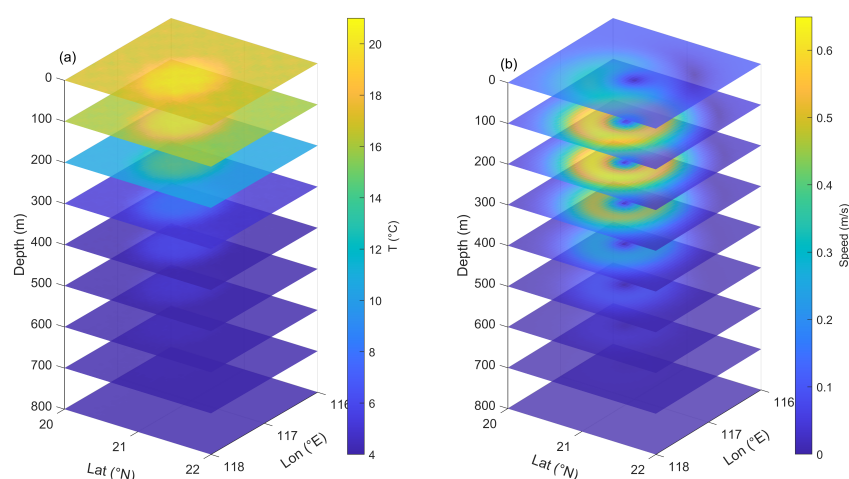


Figure 2. Three-dimensional structure of the parameterized eddy model. (a) Horizontal slices of the temperature field at 100 m intervals from the surface to 1000 m depth. (b) Horizontal slices of the Flow field at the same depth levels.

The environmental model employs an idealized axisymmetric Lamb–Oseen vortex and a parameterized temperature field, which do not capture azimuthal asymmetry, non-uniform background currents, or multi-eddy flows present in real ocean environments. These simplifications enable controlled uncertainty propagation at computational scales suitable for 5000-sample LHS ensembles. Their implications for result generality are discussed in Section 5.

2.3. Uncertainty-Integrated Coordinated Motion Model

The coordinated motion model simulates fleet trajectories under a single design principle: no adaptive control or real-time optimization is applied, so observation quality reflects only the propagation of environmental and instrumental uncertainties. The model comprises two coupled layers: an upper mission-planning layer (Section 2.3.1) and a

lower motion-integration layer (Section 2.3.2). Uncertain parameters from both layers are combined in an integrated kinematics model that propagates the fleet state through discrete-time stepping (Section 2.3.3).

2.3.1. Upper Layer: Mission Planning

The planning layer generates fleet configuration and sampling waypoints for the motion-integration layer. An edge-center architecture (Figure 3a) is chosen to support comprehensive multi-dimensional evaluation. Edge circumnavigation resolves boundary shape, which is essential for tracking eddy evolution [15,36–39]. Center transection captures radial cross-sections of temperature and thermocline structure [18,40]. Both modes have been independently validated in prior studies as effective strategies for eddy observation. Together, they cover horizontal and vertical reconstruction, thermocline characterization, and eddy center localization, ensuring that all quality dimensions are observable.

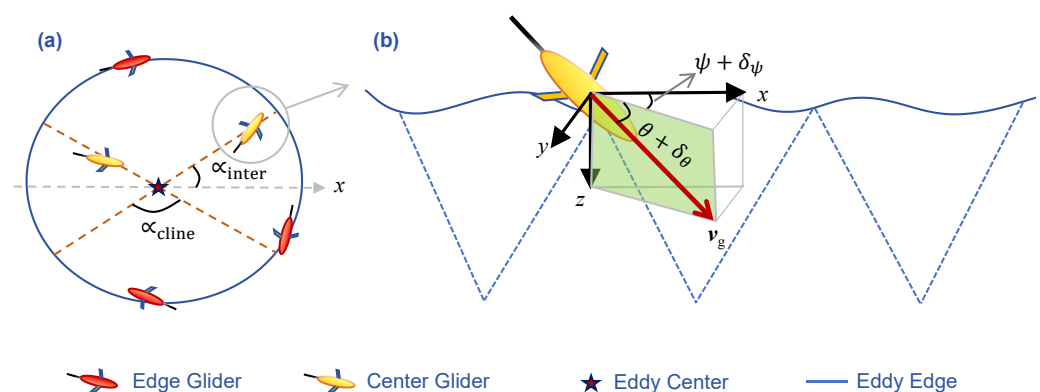


Figure 3. Fleet architecture and glider kinematics. (a) Edge-center architecture (three edge gliders, two center gliders) with centerline angle α_{cline} and intersection angle α_{inter} . (b) Glider motion parameters: pitch θ with uncertain bias, heading ψ with uncertain bias, and speed v_g .

Based on the eddy dimensions from the parameterized environmental field (Section 2.2), an architecture with three edge gliders and two center gliders (Figure 3a) is designed under the principle of spatiotemporal synchronization to reduce spatiotemporal aliasing during observation [10] and to provide a baseline for evaluating synchronization quality under uncertainty.

The two center gliders follow diametrical transects parameterized by the centerline angle α_{cline} and intersection angle α_{inter} . A total of n_c sampling waypoints are distributed uniformly along the two transects. These three parameters (α_{cline} , α_{inter} , n_c) are treated as controllable uncertain inputs in the LHS experiment, enabling analysis of how mission planning choices propagate to observation quality.

The three edge gliders equally partition the eddy perimeter and sample along the boundary. Their waypoints are determined by a kriging-variance criterion based on equal-angle sectors:

$$n_i = \text{round} \left(\frac{L_i}{\Delta s} \cdot \left[1 + \lambda_K \cdot \frac{\sigma_{K,i}^2 - \bar{\sigma}_K^2}{s_{\sigma^2}} \right] \right) \quad (5)$$

where L_i is the arc length of sector i ($i = 1, \dots, 12$), Δs is the base profile spacing, $\sigma_{K,i}^2$ is the kriging variance of sector i , $\bar{\sigma}_K^2$ is the mean variance across all sectors, s_{σ^2} is the standard deviation of the sector variances, and $\lambda_K = 0.5$ controls the sensitivity to variance. Spacing is constrained to [5, 20] km.

2.3.2. Lower Layer: Motion Integration

The SCOPE framework analyzes mission-level uncertainty propagation rather than vehicle-level dynamics. A first-order kinematic description is therefore adopted, following standard practice in glider mission simulation and path-planning studies [23,24]. This avoids platform-specific hydrodynamic parameters whose trajectory-level effects are already captured by the kinematic parameters and eliminates computational overhead with less corresponding benefit for mission-level evaluation.

In three-dimensional space, the glider moves at speed v_g in a sawtooth pattern between the surface and maximum depth (Figure 3b). Neglecting the angle of attack, the glide angle is represented by the pitch angle θ , and the horizontal heading by the heading angle ψ . The velocity components are

$$\dot{x} = v_g \cos \theta \cos \psi, \quad \dot{y} = v_g \cos \theta \sin \psi, \quad \dot{z} = v_g \sin \theta \quad (6)$$

The heading angle is updated at each surfacing event to follow the planned path. A new heading command ψ_{cmd} is computed from the current surfacing position to the next profile endpoint. A first-order lag model simulates the physical inertia of glider turning:

$$\dot{\psi} = \text{sat}\left(\frac{\psi_{\text{cmd}} - \psi}{\tau_\psi}, \dot{\psi}_{\text{max}}\right) \quad (7)$$

where $\tau_\psi = 120$ s is the time constant and $\dot{\psi}_{\text{max}} = 3^\circ/\text{s}$ is the maximum turn rate.

2.3.3. Uncertainty Integration and Propagation

The SCOPE motion model adopts an additive superposition principle: environmental disturbances and instrumental errors are superimposed onto the nominal glider kinematics. This follows the standard approach in the glider motion accuracy literature, where the total velocity is decomposed into glider through-water velocity plus the ocean current field [25,26]. The eddy current field $\mathbf{u}(\mathbf{r}, z)$ from Section 2.2 adds to the horizontal velocity components:

$$\dot{x} = v_g \cos \theta \cos \psi + u(\mathbf{r}, z), \quad \dot{y} = v_g \cos \theta \sin \psi + v(\mathbf{r}, z) \quad (8)$$

Heading and pitch angles are commanded setpoints that the glider attempts to track at each surfacing event. In practice, compass calibration errors, mounting misalignments, and current-induced heading drift cause systematic deviations from the commanded values. Following the additive bias convention established in the glider dynamics literature [25], these deviations are modeled as additive biases (δ_ψ and δ_θ) applied at each surfacing event:

$$\psi \leftarrow \psi + \delta_\psi, \quad \theta \leftarrow \theta + \delta_\theta \quad (9)$$

Gliding speed v_g is an operational variable that the control system selects from a feasible range based on mission requirements and environmental conditions. Sampling v_g uniformly within this range via LHS captures both control-level decision variability (different conditions call for different optimal speeds) and execution-level deviations (discrepancies between commanded and achieved speed). A separate additive error term δ_v is therefore unnecessary—the range already absorbs both sources of variation.

The uncertainty propagation is performed by the motion model under 10-dimensional uncertain inputs consisting of four environmental parameters, the three motion parameters and three planning parameters (Table 2). The environmental parameters V_{max} , U_c , θ_c , z_m , the planning parameters α_{inter} , α_{cline} , n_c and motion parameter v_g are sampled uniformly within ranges derived from observational data or engineering constraints. The motion

parameters δ_ψ , δ_θ follow the additive convention as additive errors sampled in each range. Table 2 provides the numerical ranges for all ten parameters. Feeding these parameters into the coordinated motion model yields the fleet sampling plan and glider trajectories, as illustrated in Figure 4 under mean parameter values.

Table 2. Input parameters of the SCOPE framework.

Symbol	Name	Range	Source
$V_{\max}$	Maximum eddy velocity	[0.2, 1.5] m/s	Environment
U_c	Background current speed	[0, 0.5] m/s	Environment
θ_c	Background current heading	[0°, 360°]	Environment
z_m	Depth of max velocity	[50, 300] m	Environment
v_g	Glider speed	[0.20, 0.80] m/s	Motion
δ_ψ	Heading bias	[−5°, 5°]	Motion
δ_θ	Pitch bias	[−3°, 3°]	Motion
α_{inter}	Intersection angle	[0°, 180°]	Planning
α_{cline}	Centerline angle	[0°, 360°]	Planning
n_c	Center sampling points	[5, 30]	Planning

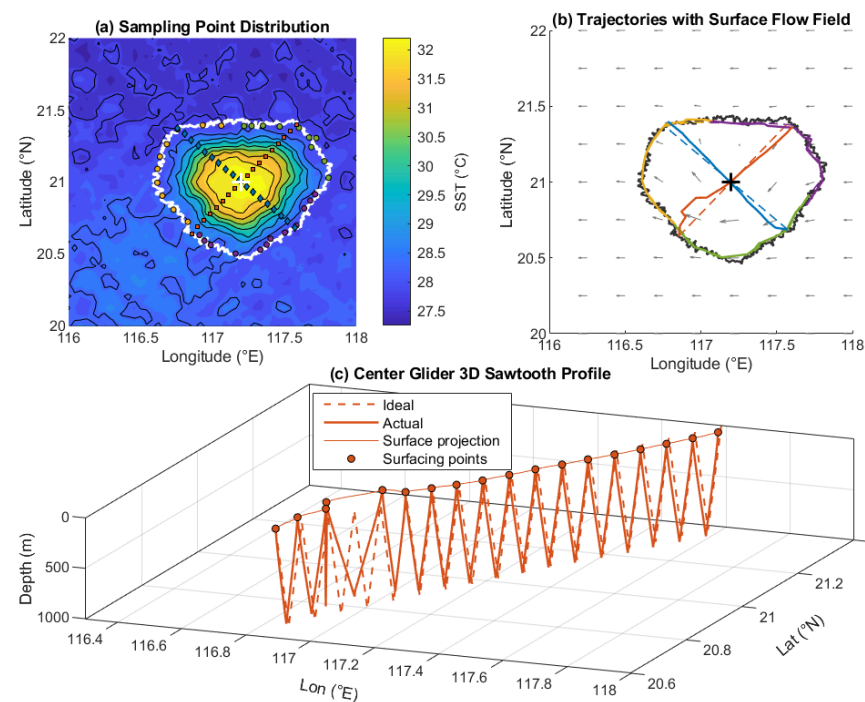


Figure 4. Trajectories under uncertainty. (a) Planned waypoints of the five-glider fleet. (b) Realized vs. planned trajectories under one uncertainty draw: dashed lines represent ideal trajectories, colored solid lines represent actual glider trajectories for each of the five gliders, thick solid line represents the eddy boundary, and gray arrows represent the flow field. (c) A 3D sawtooth profile of a center glider (theoretical vs. actual).

2.4. Performance Assessment and Core Metric Selection

An LHS [41] sample matrix is generated with $N = 5000$ draws in MATLAB R2022b (MathWorks, Natick, MA, USA). Each draw simulates a complete mission and computes 27 evaluation metrics, producing the response matrix $\mathbf{Y} \in \mathbb{R}^{5000 \times 27}$. A composite score is first constructed as a scalar measure of overall observation quality. The ACMS algorithm then compresses the metric set to a compact core subset that preserves the score-based ranking fidelity.

2.4.1. Multi-Dimensional Metric System and Composite Score

Seven quality categories of metrics are designed to span the complete observation quality chain, from individual trajectory fidelity to mission-level operational sustainability (see Supplementary Material, Section S1 for full derivations):

1. Trajectory error captures how environmental and instrumental uncertainties displace individual glider paths from their planned tracks. This is the most fundamental uncertainty manifestation—all downstream quality dimensions depend on where the gliders actually sample. Metrics include net swing error, mean and maximum absolute error, path length error, and time coefficient of variation.
2. Synchronization quantifies how consistently the fleet maintains its intended spatiotemporal formation. Poor synchronization introduces aliasing artifacts when merging asynchronous profiles into a unified eddy snapshot, degrading the interpretability of reconstructed fields. Metrics include glider-to-glider time dispersion and profile rate dispersion.
3. Eddy center localization evaluates the accuracy with which the fleet determines the eddy center position, a prerequisite for interpreting radial structure, computing azimuthal averages, and comparing with satellite-derived center estimates. Metrics include center error (Euclidean distance) and its temperature-weighted variant.
4. Boundary shape fidelity measures how well the edge-glider sampling resolves the eddy perimeter geometry. Accurate boundary delineation is essential for tracking eddy evolution, estimating eddy area and volume, and computing boundary-integrated fluxes. Metrics include area error, perimeter error, Procrustes rotation angle, and Hausdorff distance.
5. Horizontal temperature mapping assesses the accuracy with which the reconstructed surface and subsurface temperature fields reproduce the ground truth. This reflects the core oceanographic objective of glider-based eddy observation: resolving the three-dimensional thermal structure. Metrics include surface max error, surface RMSE, subsurface max error, and subsurface RMSE.
6. Vertical section profiling evaluates how well the center-glider transects resolve vertical thermal stratification and thermocline structure. Accurate vertical resolution is critical for computing heat content, available potential energy, and eddy-induced transport. Metrics include thermocline depth RMSE, thermocline strength error, vertical max error, and vertical grid coverage.
7. Energy efficiency addresses the operational sustainability dimension: given limited onboard battery capacity, how efficiently does the fleet convert energy expenditure into observation quality? This is a first-order practical constraint for multi-day eddy surveys. Metrics include energy per unit distance, energy per profile, battery utilization balance, and depth efficiency.

For each of the 5000 LHS draws, the 10 uncertain parameters perturb the glider trajectories through the motion model, and the 27 metrics are computed from the resulting sampling outcome. Each metric is normalized to $[0, 1]$ according to its directionality, with bounds fixed from the 5000-sample ensemble. The composite score is the equally weighted mean of the 27 normalized metrics, serving as the reference for metric selection. The robustness of equal weighting against alternative schemes (entropy-weight, category-weight, CV-weight, and random) is validated in Supplementary Material, Section S2.

2.4.2. Adaptive Core Metric Selection

The ACMS algorithm is designed to reduce redundancies among metrics while maximally preserving multi-dimensional evaluation information. Unlike conventional feature selection methods (PCA, mRMR, recursive feature elimination, mutual information se-

lection, and LASSO) that optimize for prediction accuracy, ACMS optimizes ranking fidelity—the Spearman correlation between the reduced metric set’s composite score and the full 27-metric reference score—because the goal is to preserve the relative ordering of uncertainty scenarios rather than to predict a specific output value. It also enforces a category-coverage constraint, ensuring each of the seven quality dimensions retains at least one representative regardless of its variance magnitude. Starting from all 27 metrics, ACMS iteratively removes the metric whose elimination causes the smallest drop in ρ_S , subject to the constraint that each category retains at least one representative:

$$\mathcal{M}^* = \arg \min_{\mathcal{M}} |\mathcal{M}| \quad \text{s.t.} \quad \rho_S(\mathbf{S}(\mathcal{M}), \mathbf{S}_{\text{ref}}) \geq 0.95 \quad (10)$$

where $\mathcal{M}^*$ denotes the optimal core metric set, $\mathcal{M}$ denotes a candidate metric set, $|\mathcal{M}|$ is its cardinality, $\mathbf{S}(\mathcal{M})$ is the composite score computed from $\mathcal{M}$, and $\mathbf{S}_{\text{ref}}$ is the reference score from the full 27-metric set.

ACMS operates in three phases (full pseudo-code in Supplementary Material, Section S3). Phase 1 pre-filters by coefficient of variation ($\text{CV} > 0.10$) and enforces category coverage. Phase 2 iteratively removes the metric whose elimination causes the smallest drop in Spearman correlation ρ_S with the reference composite score, stopping when $\rho_S < 0.95$ or further removal would empty a category. Phase 3 computes the final composite score from the retained metrics.

Eighteen elimination steps yield $|\mathcal{M}^*| = 9$ core metrics with $\rho_S = 0.967$ (Figure 5), a 67% reduction while preserving all seven categories. The nine core metrics are: net swing error, time coefficient of variation, temperature-weighted center error, Procrustes rotation angle, thermocline RMSE, surface max error, vertical max error, vertical grid coverage, and energy per unit distance. Five-fold cross-validation confirms no overfitting (details in Supplementary Material, Section S4). The 9-metric composite score is used for all subsequent analyses.

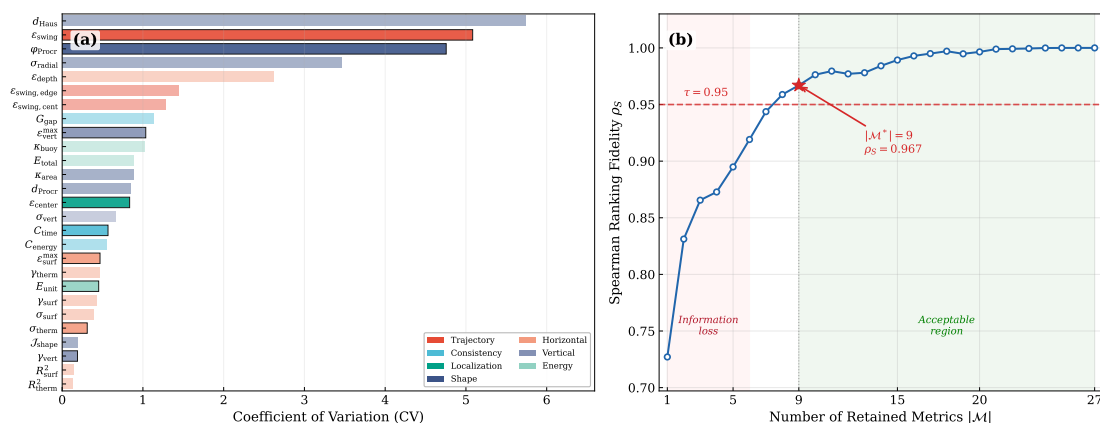


Figure 5. ACMS metric selection. (a) Coefficient of variation (CV) of all metrics. (b) Spearman correlation ρ_S during backward elimination; the star marks the 9-metric solution ($\rho_S = 0.967$, $\tau = 0.95$ dashed). The 8-metric solution is rejected for failure in all quality categories coverage.

3. Sensitivity Analysis Under Uncertainty

The effects of the 10 uncertain parameters on observation quality are analyzed in this section. First, correlation and interaction analysis is performed on the original parameters to rank their importance and identify physical couplings. Second, derived variables are constructed to resolve these couplings, followed by dose–response characterization of the dominant parameters. All results are based on the 5000-sample LHS ensemble and the 9-metric composite score from Section 2.4.

3.1. Correlation and Interaction Analysis

Parameter importance is assessed using Spearman's rank correlation ρ_S and k -NN mutual information (MI, $k = 5$) [42,43]. With $N = 5000$ samples, even small effects achieve nominal significance. Practical significance is therefore assessed using Cohen's d [44], with results reported as meaningful only when $p < 0.05$ and $|d| \geq 0.10$ [45]. As shown in Figure 6, $V_{\max}$ and v_g are the top-ranked parameters, suggesting a strong coupling between them. This motivates the interaction analysis in Section 3.2.

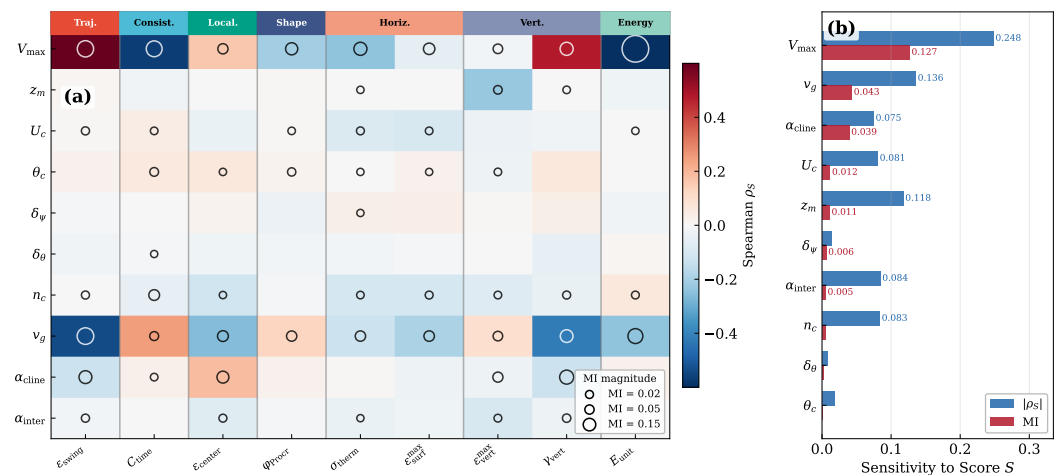


Figure 6. Two-stage sensitivity analysis. (a) Spearman ρ_S (color) and MI (circle size) of the 10 raw parameters against the 9 core metrics. (b) $|\rho_S|$ and MI of each parameter against the composite score, ordered by MI.

3.2. Interaction Analysis and Derived Variables

To resolve the couplings suggested by the correlation analysis, all $\binom{10}{2} = 45$ pairs of the 10 input parameters are assessed using factorial ANOVA with Bonferroni correction ($\alpha^* = 0.05/45 = 0.0011$). Three pairs exhibit notable interactions (Figure 7): $V_{\max} \times v_g$ ($\eta^2 = 0.035$, partial eta-squared, ANOVA effect size), $U_c \times v_g$ ($\eta^2 = 0.010$), and $\theta_c \times \alpha_{\text{cline}}$ ($\eta^2 = 0.007$). The $V_{\max} \times v_g$ interaction is the strongest, confirming that glider speed and eddy intensity share a coupled physical mechanism. The $\theta_c \times \alpha_{\text{cline}}$ interaction reveals that performance depends on the relative angle between background current and transect heading. The $U_c \times v_g$ interaction ($\eta^2 = 0.010$) does not admit a simple derived variable, and U_c is rarely identifiable in real ocean environments, so this coupling is not further resolved. The remaining 93% of interactions have $\eta^2 < 0.005$, indicating that the system is approximately additive once the dominant couplings are resolved. These couplings motivate the construction of physically independent derived variables. The velocity ratio

$$\eta_v = \frac{v_g}{V_{\max}} \quad (11)$$

absorbs the $v_g \times V_{\max}$ interaction into a single dimensionless parameter. The background current is decomposed relative to the transect heading:

$$U_c^\perp = U_c \sin(\theta_c - \alpha_{\text{cline}}), \quad U_c^\parallel = U_c \cos(\theta_c - \alpha_{\text{cline}}) \quad (12)$$

where $U_c^\perp$ measures lateral drift and $U_c^\parallel$ measures headwind/tailwind effects [42], resolving the $\theta_c \times \alpha_{\text{cline}}$ coupling into physically interpretable components.

The sensitivity analysis is repeated with the derived variables using the same methods. η_v ranks first by MI (0.375). Its low $|\rho_S| = 0.141$ relative to the high MI confirms a non-monotonic effect: both excessively low and excessively high η_v degrade performance.

The dose–response characteristics of η_v are examined in Section 3.3 to characterize this relationship. $U_c^\perp$ ranks second among environmental factors ($|\rho_S| = 0.163$, $MI = 0.037$), indicating that lateral drift is the dominant pathway through which background currents affect observation quality, while the along-track component $U_c^\parallel$ shows negligible influence ($|\rho_S| = 0.004$). Table 3 compares the two parameter groups, providing a ranked reference for mission design. Parameters with $p \geq 0.05$ are marked as “Not sig” (not statistically significant).

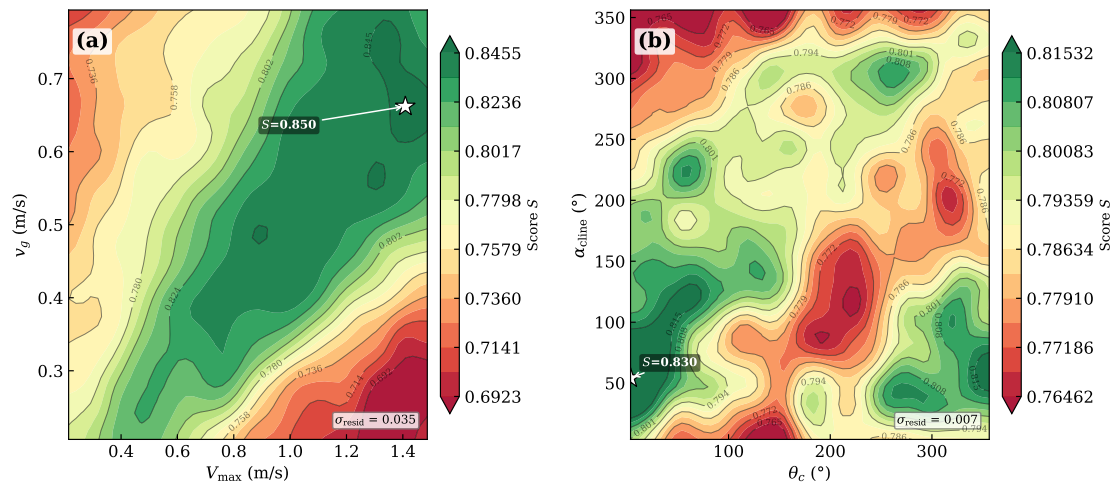


Figure 7. Interaction effects between key parameter pairs. (a) $V_{\max}$ – v_g interaction ($\eta^2 = 0.035$), the strongest among all 45 pairs. (b) θ_c – α_{cline} interaction ($\eta^2 = 0.007$). Stars mark the parameter combination yielding the highest composite score in each contour plot.

Table 3. Parameter importance ranking from 5000-sample LHS ensemble. The 10 original parameters vs 9 derived parameters after decoupling $v_g \times V_{\max}$ and $\theta_c \times \alpha_{\text{cline}}$.

Rank	Parameter	ρ_S	p	MI	MI Rank	Type
<i>Original parameters (10)</i>						
1	$V_{\max}$	+0.248	10^{-70}	0.111	1	Environmental
2	v_g	+0.136	10^{-22}	0.043	2	Controllable
3	z_m	+0.118	10^{-17}	<0.001	9	Environmental
4	α_{inter}	+0.084	10^{-9}	0.005	8	Controllable
5	n_c	+0.084	10^{-9}	0.009	6	Controllable
6	U_c	+0.081	10^{-8}	0.012	5	Environmental
7	α_{cline}	−0.075	10^{-7}	0.039	3	Controllable
8	θ_c	−0.018	0.21	<0.001	10	Not sig
9	δ_ψ	−0.014	0.34	0.020	4	Not sig
10	δ_θ	+0.008	0.59	0.002	7	Not sig
<i>Derived parameters (9)—after decoupling</i>						
1	η_v	−0.141	10^{-23}	0.375	1	Controllable
2	$V_{\max}$	+0.248	10^{-70}	0.127	2	Environmental
3	$U_c^\perp$	−0.163	10^{-31}	0.037	3	Environmental
4	z_m	+0.118	10^{-17}	0.011	5	Environmental
5	ξ_I	−0.084	10^{-9}	0.005	7	Controllable
6	n_c	+0.084	10^{-9}	0.005	8	Controllable
7	$U_c^\parallel$	+0.004	0.81	0.018	4	Environmental
8	δ_ψ	−0.014	0.34	0.006	6	Not sig
9	δ_θ	+0.008	0.59	0.002	9	Not sig

3.3. Dose–Response Characterization

The dominance of η_v motivates a detailed examination of its effect on observation quality. The 5000-sample LHS ensemble is stratified by η_v , and the mean composite score is computed for each bin.

The dose–response curve reveals a non-monotonic relationship with an optimal plateau (Figure 8a). The peak occurs near $\eta_v \approx 0.5$, with the 95%-of-peak interval spanning $[0.40, 0.76]$. This wide plateau indicates that moderate deviations from the optimal value incur limited performance penalties, providing practical tolerance for speed selection under imperfect environmental knowledge. Score distributions shown in Figure 8b are tightest near the optimal η_v and widen toward both extremes, indicating that the optimal configuration is robust to small speed variations.

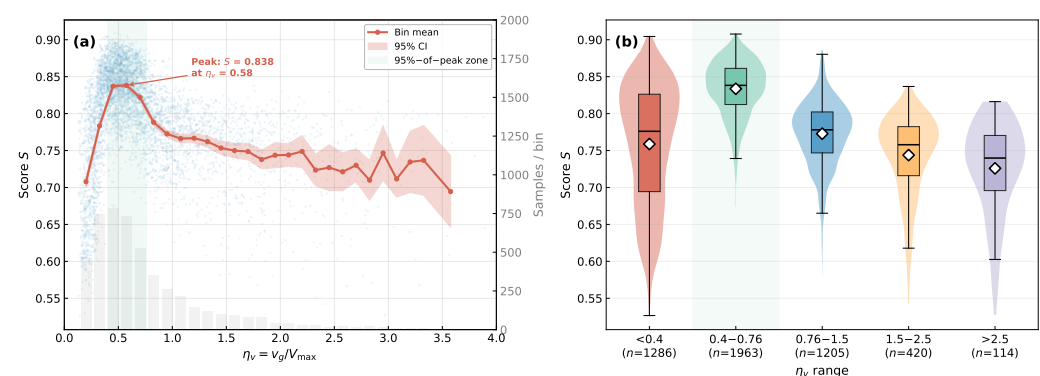


Figure 8. Dose–response of the composite score to η_v . (a) Scatter plot with binned means and 95% confidence intervals; the green region marks the 95%-of-peak interval. (b) Score distributions for discrete η_v ranges.

Since the dose–response curve exhibits attenuation at both low and high η_v , the LHS samples are further stratified by V_{max} and U_c to examine whether the optimal η_v depends on eddy intensity. Figure 9a reveals a systematic shift: the optimal η_v decreases monotonically with V_{max} (from ~ 0.58 for weak eddies to ~ 0.47 for strong eddies, Pearson $r = -0.87$), and the optimal interval narrows. The overall pattern is that weak eddies favor higher η_v to ensure sufficient spatial coverage, while strong eddies benefit from lower η_v , allowing the glider to move with the flow and accumulate more profiles within the high-velocity region. Figure 9b shows that the optimal η_v is largely insensitive to U_c , as the dose–response curves for different background current magnitudes nearly overlap. This confirms that V_{max} , not U_c , is the primary environmental driver of η_v selection. This trend is further examined on real eddies in Section 4.2.

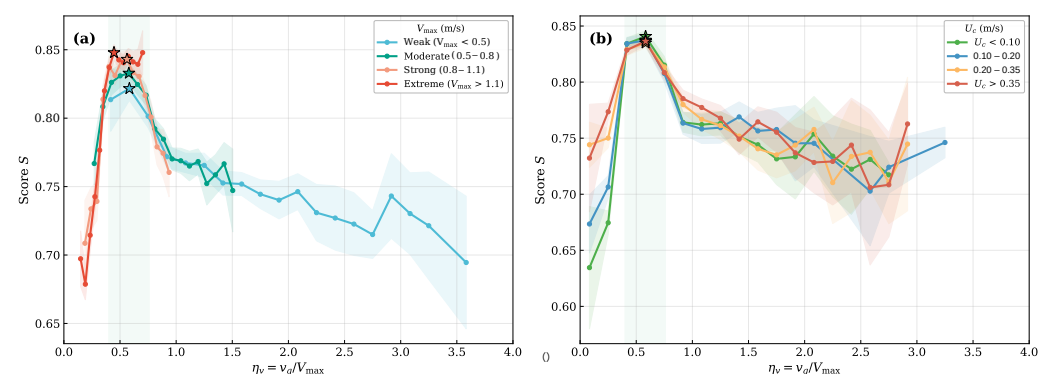


Figure 9. Intensity-dependent η_v dose–response. (a) Curves stratified by V_{max} group, showing the peak shifting from ~ 0.58 (weak) to ~ 0.47 (extreme). (b) Stratified by background current magnitude U_c .

3.4. Ablation Validation

The sensitivity analysis above identifies two actionable design rules: (i) match η_v to eddy intensity within the optimal plateau, and (ii) align α_{cline} perpendicular to the background current direction. To test whether these rules causally improve observation quality rather than merely correlate with it, a structured ablation experiment is conducted under a fixed environment ($V_{\text{max}} = 0.85$ m/s, $U_c = 0.25$ m/s). Five groups are compared under $N = 200$ paired uncertainty draws (Table 4): A (deterministic ceiling via grid search), representing the optimal design under certainty; B (literature baseline, $v_g = 0.30$ m/s) [10], representing the degradation of A under uncertainty; C (η_v rule, v_g sampled from $\eta_v \sim \mathcal{U}[0.40, 0.60] \times V_{\text{max}}$, a suitable interval for the moderate eddy intensity used in this experiment, Section 3.3); D ($\alpha_{\text{cline}} = \theta_c + 90^\circ$), testing the alignment rule identified in Section 3.2; and E (C and D combined). All groups share identical secondary parameters from Group A. Paired differences are assessed using the Wilcoxon signed-rank test [46]. The Normalized Gap Closure NGC = $(\bar{S}_{\text{group}} - \bar{S}_B) / (\bar{S}_A - \bar{S}_B) \times 100\%$ quantifies the fraction of the A–B gap recovered by each strategy [47].

Figure 10a shows the score distributions. All group means lie below the deterministic ceiling, and the ceiling falls within the upper tail of the distributions, indicating that the grid search may not have fully captured the parameter space. Group B exhibits the largest variance, confirming the baseline strategy as the least robust. Groups C–E all show markedly reduced variance. Figure 10b,c show the paired differences and NGC values (Table 4).

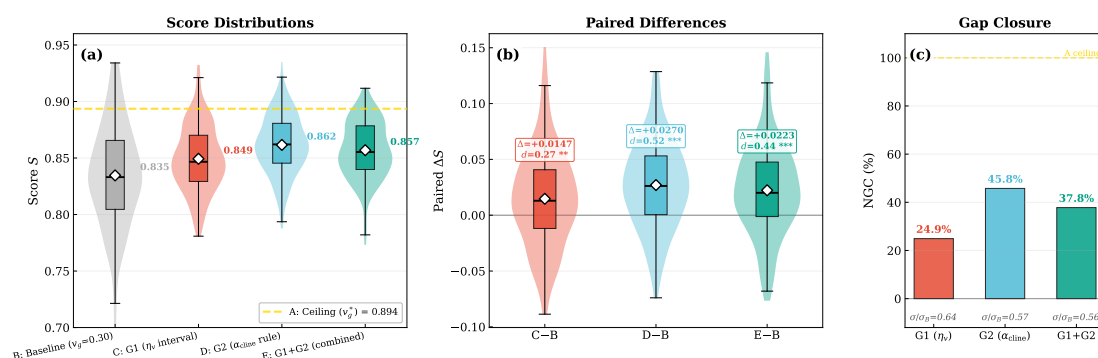


Figure 10. Ablation results ($N = 200$ paired draws). (a) Score distributions for Groups A–E; the yellow line marks the deterministic ceiling. (b) Paired score differences relative to B. Stars indicate Cohen’s d effect size: *** $d \geq 0.50$, ** $d \geq 0.25$. (c) Normalized Gap Closure.

Table 4. Ablation results ($N = 200$ paired samples, $V_{\text{max}} = 0.85$ m/s, $U_c = 0.25$ m/s). Groups D and E use $\alpha_{\text{cline}} = \theta_c + 90^\circ$; Groups C and E sample η_v from $[0.40, 0.60]$.

Group	$\bar{S}$	Std	Δ vs. B	d	p	NGC
A (ceiling)	0.894	—	—	—	—	100%
B (literature)	0.835	0.046	0	—	—	0%
C (η_v interval)	0.849	0.029	+0.015	0.27	4.0×10^{-3}	25.0%
D (α_{cline} rule)	0.862	0.026	+0.027	0.52	$<10^{-11}$	45.8%
E (combined)	0.857	0.025	+0.022	0.44	2.0×10^{-8}	37.8%

The A $\rightarrow$ B gap (0.059) quantifies a 6.6% performance penalty for using a fixed literature-default speed. While the absolute score differences are modest on the $[0, 1]$ scale, the NGC metric normalizes this gap relative to the achievable improvement, isolating the contribution of each design rule. Whether a 6.6% improvement translates to operationally meaningful gains depends on the specific mission’s tolerance for reconstruction error. The α_{cline} rule alone recovers 45.8% of this gap ($d = 0.52$, $p < 10^{-11}$; effect sizes by Cohen’s

d), outperforming the η_v rule (25.0%, $d = 0.27$). The combined strategy (37.8%) shows mild negative synergy, demonstrating the framework's ability to detect suboptimal parameter interactions. All rule-based strategies reduce performance variability by 36–44%, providing downside protection independent of mean improvement.

The ablation yields three practical insights. First, transect orientation is a high-leverage design choice, recovering nearly half of the performance gap. Although perpendicular alignment maximizes lateral drift, the drift disperses sampling points over a wider area, improving spatial coverage and temperature reconstruction; the composite score registers a net gain because these coverage benefits outweigh the associated trajectory penalties. Second, the η_v rule provides a smaller but consistent improvement, robust to the specific choice of interval within the optimal plateau. Third, the combined strategy does not add the two improvements, revealing a synergy structure otherwise invisible to single-factor analysis—an advantage of the framework's ability to decompose joint effects.

4. Methodology Transfer to Real Ocean Eddies

The analyses in Section 3 are based on the idealized temperature model and Lamb–Oseen eddy model. These analytical simplifications—axisymmetric vortex geometry, uniform background current, and a parameterized temperature anomaly—enable controlled uncertainty propagation but may produce findings that differ quantitatively from real ocean conditions. To evaluate whether the SCOPE framework's core methodology (ACMS metric selection, parameter ranking, and η_v dose–response characterization) transfers to non-idealized environments, four real mesoscale eddies in the South China Sea (Figure 11) are extracted from the HYCOM GLBy0.08 reanalysis product [48,49] (1/12° resolution). Table 5 summarizes their properties.

The eddies span $V_{\max} = 0.28$ – 0.98 m/s, covering weak to strong regimes. Eddy boundaries are identified using a multi-criterion consensus method combining the Okubo–Weiss parameter, stream function closed contours, and velocity e -folding contour (see Supplementary Material, Section S5 for computational details).

Table 5. Properties of four real South China Sea eddies.

Eddy	Period	$V_{\max}$ (m/s)	$R_{\max}$ (km)	Ellipticity
E1 (weak)	2017 Summer	0.28	79	1.11
E2 (mod-weak)	2020 Summer	0.39	76	1.22
E3 (moderate)	2019 Summer	0.51	64	1.28
E4 (strong)	2020 Winter	0.98	104	1.15

The actual 3D velocity field from HYCOM replaces the analytical model in the glider motion simulator, while all other modules remain unchanged. The Lamb–Oseen radial profile correlates with the real eddy velocity fields at a mean Pearson $r = 0.918$ across the four eddies (computed on azimuthally averaged radial profiles), and the radial temperature anomaly profile correlates at $r = 0.981$, confirming that the analytical model captures the essential velocity and temperature structure. Three eddy centers are shown in Figure 11: the temperature-derived center (red star), the vorticity-derived center (white cross), and the surface cold-water center (green circle) of E4, which is a surface-cold anticyclonic eddy. The temperature-derived center is used as the reference for eddy center localization metrics (Section 2.4).

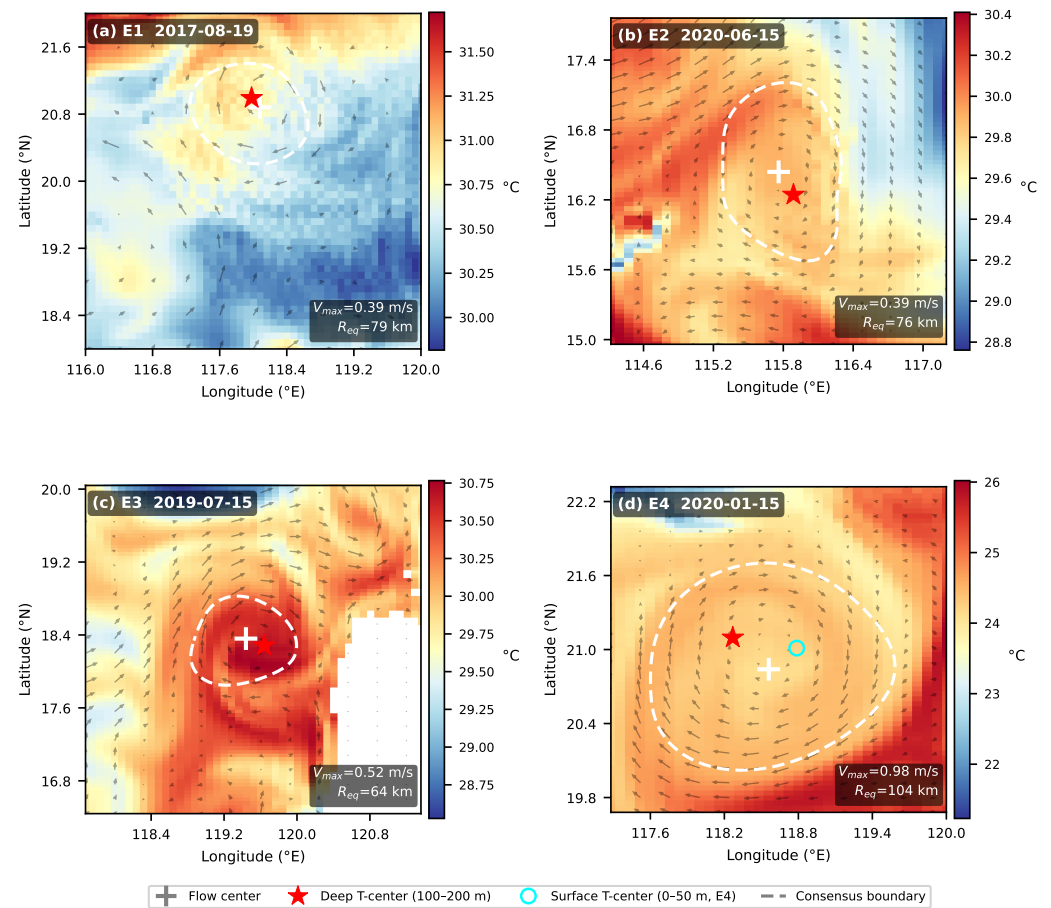


Figure 11. Real eddy velocity fields and consensus boundaries (white solid). Surface current vectors (arrows) overlaid on SST. Red star: temperature center; white cross: vorticity center; green circle (E4): surface cold center.

4.1. ACMS Framework Transferability and Parameter Ranking

The ACMS algorithm is independently applied to each real eddy to test whether the SCOPE assessment framework operates meaningfully on non-idealized data. For each eddy, 500 LHS simulations using the actual HYCOM 3D velocity field yield a response matrix $\mathbf{Y} \in \mathbb{R}^{500 \times 27}$, with the same 27 candidate metrics used in the analytical model. All four eddies achieve $\rho_S > 0.95$ (Table 6), confirming reliable convergence on non-idealized data. We note that the 500-sample ensemble per real eddy, while sufficient for Spearman correlation estimation (statistical power > 0.80 for $|\rho_S| \geq 0.12$), may yield conservative mutual information estimates for weak parameter dependencies. The consistent ranking of η_v across all four eddies and its agreement with the analytical 5000-sample result (MI rank = 1) mitigate this concern.

Table 6. ACMS backward elimination results for each real eddy, compared with the analytical model. $|\mathcal{M}^*|$: number of core metrics retained; ρ_S : Spearman correlation between core and full composite scores; Overlap: number of exact matches with the 9 analytical core metrics (Sim-9).

Environment	$ \mathcal{M}^* $	ρ_S	Overlap	Recall
Analytical model	9	0.967	—	—
E1 (Weak, $V_{\max} = 0.28$ m/s)	13	0.955	4/9	44%
E2 (Mod-Weak, $V_{\max} = 0.39$ m/s)	14	0.952	5/9	56%
E3 (Moderate, $V_{\max} = 0.51$ m/s)	20	0.955	8/9	89%
E4 (Strong, $V_{\max} = 0.98$ m/s)	15	0.956	6/9	67%

The number of retained core metrics ranges from 13 to 20, larger than the analytical model's 9, reflecting the richer variability structure of real eddies: asymmetric temperature fields and complex flows activate additional quality dimensions. The variation in $|\mathcal{M}^*|$ across eddies reveals a systematic relationship with eddy intensity. E3 (moderate, $V_{\max} = 0.51$ m/s) retains the largest set (20 metrics) because its velocity field lies within the analytical model's calibrated range and its intermediate intensity yields sufficient performance variation to make many metrics discriminable. E1 (weak, $V_{\max} = 0.28$ m/s) retains only 13 metrics because the weaker velocity field reduces the dynamic range of trajectory and shape metrics, pushing more metrics below the CV threshold. E4 (strong) occupies an intermediate position (15 metrics) because its intense velocity field amplifies trajectory errors but suppresses the discriminability of energy metrics—the buoyancy subsystem dominates energy consumption regardless of parameter choices. This intensity-dependent metric activation is a feature of the framework rather than a shortcoming: it reveals which quality dimensions are practically monitorable for a given eddy regime, guiding mission designers toward a leaner, regime-appropriate evaluation set.

Among the 9 analytical core metrics, 5 are stably reproduced across ≥ 3 of 4 eddies: vertical max error and vertical coverage (4/4), Procrustes rotation, thermocline RMSE, and energy per km (3/4). E3 ($V_{\max} = 0.51$ m/s, closest to the analytical model's range) achieves 89% overlap (8/9). The complete per-eddy core metric lists are provided in the Supplementary Material, Section S6.

The rankings of the uncertain parameters in real eddies are then analyzed using the method proposed in Section 3.1. Cross-environment consistency is assessed using Kendall's coefficient of concordance [50]:

$$W = \frac{12 \sum_{j=1}^k (R_j - \bar{R})^2}{p^2(k^3 - k)} \quad (13)$$

where R_j is the rank-sum for parameter j across $p = 4$ judges (the four real eddies) and $k = 6$ parameters. The result $W = 0.593$ ($\chi^2 = 11.86$, $p = 0.037$) indicates statistically significant agreement among the four real eddies ($p < 0.05$), ruling out random coincidence in cross-eddy parameter ranking consistency. We note that the observed concordance is moderate and the p -value is near the conventional threshold of 0.05. This reflects genuine variability in parameter importance across eddy regimes rather than a failure of the methodology. For instance, α_{cline} ranks second overall but drops to fifth in E4—the strong, near-axisymmetric eddy where centerline orientation matters less because the velocity field is more radially symmetric. Caution is therefore warranted when extrapolating parameter rankings from one eddy regime to another; the framework identifies which parameters are consistently important across regimes (η_v) and which are regime-dependent (α_{cline} , n_c).

The global ranking is $\eta_v > \alpha_{\text{cline}} > n_c > \delta_\theta > \delta_\psi > \alpha_{\text{inter}}$, with η_v ranking first in 2 of 4 eddies and within the top 3 in all cases. Including the analytical model as a fifth judge yields $W = 0.515$ ($p = 0.025$). The dominant parameter (η_v) and the least influential parameters (α_{inter} , δ_ψ) are consistent across environments, while intermediate rankings vary with eddy intensity. The per-eddy sensitivity rankings are documented in the Supplementary Material, Section S6.

4.2. η_v Dose–Response Transferability

The intensity-dependent η_v dose–response observed in the analytical model (Section 3.3) is examined in the real eddies to investigate the systematic shift of the optimal η_v with $V_{\max}$. The same binned method is applied to the 500-sample LHS ensemble of each real eddy.

Figure 12 summarizes the results. The optimal η_v —marked by the peak of each dose–response curve—decreases monotonically with eddy intensity in both datasets.

For the analytical model, the peak shifts from $\eta_v \approx 0.58$ (weak eddies) to $\eta_v \approx 0.47$ (strong eddies), with Pearson $r = -0.87$ between η_v^{opt} and V_{max} . For the real eddies, the peak shifts from $\eta_v \approx 2.35$ (E1, $V_{\text{max}} = 0.28$ m/s) to $\eta_v \approx 0.35$ (E4, $V_{\text{max}} = 0.98$ m/s), with $r = -0.96$. The stars in Figure 12a–c mark the peak of each dose–response curve. The dashed lines in Figure 12d mark the LHS-optimal η_v from the 500-sample ensemble, confirming that the binned peaks are consistent with the empirical optima.

The decreasing trend is qualitatively consistent between simulation and real eddies: stronger eddies consistently favor lower η_v . Quantitatively, however, the real-eddy optimal η_v values are systematically higher than the analytical predictions, with the largest discrepancy for the weakest eddy (2.35 vs. 0.58 for E1). This offset reflects two differences between the idealized and real environments.

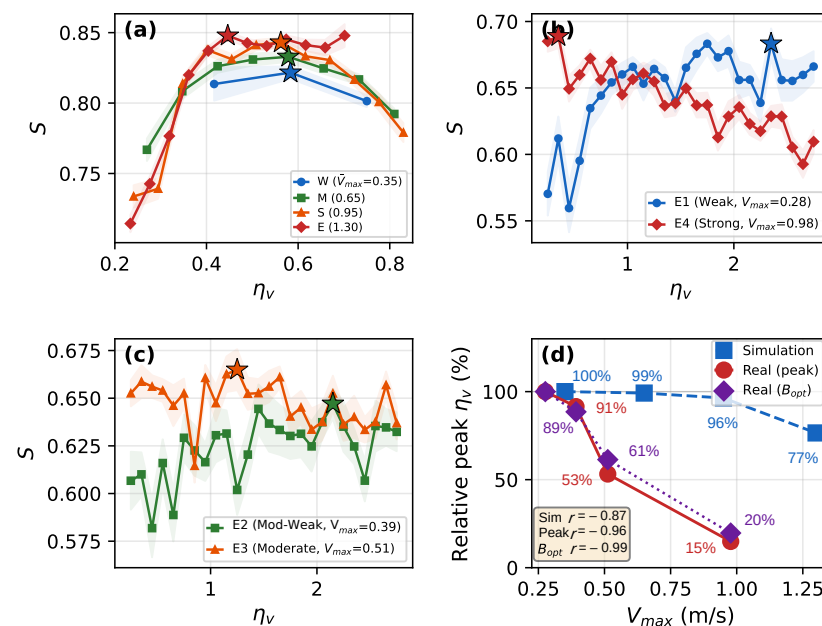


Figure 12. Optimal η_v trend: simulation vs. real eddies. (a) Analytical model: binned dose–response curves by V_{max} quartile; stars mark the peak of each curve. (b) Real eddies E1 and E4; stars mark the peak. (c) Real eddies E2 and E3; stars mark the peak. (d) η_v^{opt} vs. V_{max} for simulation (squares) and real eddies (circles); dashed lines mark the LHS-optimal η_v from the 500-sample ensemble.

First, the irregular, non-axisymmetric velocity field of real eddies induces larger and less predictable trajectory deviations than the smooth Lamb–Oseen vortex. In weak eddies, a higher glider speed enables the glider to resist lateral drift and maintain course corrections, producing a wave-like trajectory around the nominal path that increases the number of sampling points and benefits eddy center localization and temperature reconstruction. In strong eddies, by contrast, the glider has limited ability to counteract the intense flow; a high speed amplifies energy consumption and trajectory inconsistency without improving coverage. A lower speed allows the glider to move with the flow along a spiral path, accumulating more profiles within the eddy interior at lower energetic cost.

Despite the numerical offset, the monotonic decreasing trend is robust across both datasets, confirming that the intensity-dependent trend of optimal η_v transfers from the analytical model to real ocean environments. Calibrating a corrected $\eta_v^{\text{opt}}(V_{\text{max}})$ function using a broader range of real eddy data is left for future work.

5. Discussion and Conclusions

5.1. Discussion

The SCOPE framework is proposed for evaluating and diagnosing multi-glider coordinated observation of mesoscale eddies under uncertainty. Its contribution lies in integration and adaptation rather than invention of individual methods: uncertainty propagation builds on established LHS techniques [41] but extends them from single-glider trajectory errors [25,26] to fleet-level, multi-dimensional observation quality; the sensitivity analysis adopts standard correlation and interaction methods [42] but adapts them to a two-stage diagnostic pipeline with derived-variable construction for coupled parameters; the ablation validation follows the robust optimization framework [47] but applies it to decompose marginal and combined effects of observation design rules. Its core methodology includes a multi-dimensional metric system (27 metrics compressed to 9 core indicators), ACMS adaptive core metric selection, two-stage sensitivity analysis, and ablation-based diagnosis. These methods are independent of any specific simulation model and can be applied to both pre-deployment assessment with analytical environments and post-mission diagnosis with observed data.

SCOPE differs from conventional Observing System Simulation Experiments (OSSEs) in two respects. First, whereas OSSEs evaluate a proposed observing system's impact on forecast skill by assimilating synthetic observations from a single Nature Run, SCOPE does not involve data assimilation. Instead, it samples uncertain environmental parameters via LHS and evaluates how observation quality varies across an ensemble of eddy realizations, yielding a distribution of outcomes that quantifies robustness rather than a single impact score. Second, OSSEs typically assess performance through a single forecast error metric, whereas SCOPE employs a 27-metric system spanning seven quality dimensions, providing a multi-faceted characterization of observation quality. In addition, while Lermusiaux [31] established a general framework for adaptive sampling and uncertainty quantification in ocean observation, their focus was on data assimilation and adaptive path planning; SCOPE complements this by providing a diagnostic layer that decomposes performance into metric-level contributions and identifies dominant uncertainty drivers.

In the analytical model, the two-stage sensitivity analysis identifies physical couplings among parameters and reveals the non-monotonic dominant effect of η_v through derived variable construction. $U_c^\perp$ ranks second among environmental factors ($|\rho_S| = 0.163$, $MI = 0.037$), indicating that lateral drift is the dominant pathway through which background currents affect observation quality. This result is specific to the analytical model; in real eddies, the background current is spatially varying and cannot be decomposed into a single direction, so $U_c^\perp$ is not separately evaluated. The majority of parameter interactions (93%) are negligible, indicating approximate additivity once key couplings are resolved. The dose–response analysis shows that the optimal η_v decreases monotonically with $V_{\max}$: weak eddies require higher η_v , and strong eddies require lower η_v .

The ablation experiment validates the framework's diagnostic capability by decomposing marginal effects, detecting negative synergy, and distinguishing mean improvement from variance reduction. The ablation is performed exclusively on the analytical model because isolating causal effects requires a controlled environment with fixed environmental parameters and design rules as the only varying factors. For real eddies, the stratified dose–response analysis serves a parallel role: the natural variation in eddy intensity across the four HYCOM cases reproduces the monotonic $\eta_v^{\text{opt}} - V_{\max}$ trend ($r = -0.96$), confirming that the intensity dependence of optimal η_v is observable in realistic ocean fields.

The transferability of the framework's methodology is tested on four HYCOM eddies in the South China Sea. The ACMS algorithm converges on all four eddies ($\rho_S > 0.95$), and the parameter importance rankings are statistically consistent across environments

(Kendall $W = 0.59$, $p = 0.037$). The decreasing trend of optimal η_v with $V_{\max}$ is reproduced in the real eddies (Pearson $r = -0.96$), matching the analytical model ($r = -0.87$). However, the real-eddy optimal η_v values are systematically higher than the analytical predictions, particularly for weak eddies (2.35 vs. 0.58). This offset is attributed to two mechanisms. First, the irregular, non-axisymmetric velocity field of real eddies induces larger trajectory deviations than the smooth Lamb–Oseen vortex; in weak eddies, higher glider speed produces a wave-like trajectory that increases sampling density, while in strong eddies, lower speed allows the glider to move with the flow along a spiral path, accumulating more profiles at lower energetic cost. Second, the real eddies' non-axisymmetric geometry introduces localized advective surges that further increase the speed demand in weak-eddy conditions.

Four limitations should be noted. First, the analytical model uses an axisymmetric Lamb–Oseen vortex with a uniform background current. Real eddies exhibit azimuthal asymmetry, non-uniform background currents, filamentary structures, and multi-eddy interactions that are absent from the idealized model. These differences produce the systematic offset in optimal η_v between the analytical predictions and real-eddy results (Section 4.2), particularly for weak eddies where non-axisymmetric features dominate the velocity field experienced by the gliders. The transferability analysis in Section 4 partially addresses this limitation by testing on four HYCOM eddies, but broader real-ocean validation is needed. Second, a continuous function $\eta_v^{\text{opt}}(V_{\max})$ has not been derived; only discrete intensity-group trends are available, and calibration across a denser eddy intensity spectrum is needed. Third, the validation uses four eddies from a single region (South China Sea). Although these eddies span a broad intensity range ($V_{\max} = 0.28\text{--}0.98$ m/s) and the parameter rankings exhibit statistically significant cross-eddy concordance ($W = 0.593$, $p = 0.037$), four eddies constitute a small sample for generalization. Expanded multi-region validation covering a denser eddy intensity spectrum and diverse geographic conditions is needed to establish the generality of the $\eta_v^{\text{opt}}(V_{\max})$ relationship. Fourth, the 500-sample ensemble per real eddy, while adequate for rank-correlation estimation, may yield conservative MI estimates for weak dependencies; larger ensembles would improve the resolution of secondary parameter effects.

5.2. Conclusions

1. SCOPE framework. A 27-metric evaluation system spanning seven quality dimensions is constructed. The ACMS algorithm compresses the 27 metrics to 9 core indicators ($\rho_S = 0.967$) while preserving category coverage and over 95% of ranking information. The framework supports both pre-deployment assessment and post-mission diagnosis.
2. Sensitivity hierarchy and ablation diagnosis. Two-stage sensitivity analysis identifies η_v as the dominant parameter (MI = 0.375) with a non-monotonic effect; $U_c^\perp$ emerges as the leading environmental factor. The majority of parameter interactions (93%) are negligible. The optimal η_v decreases monotonically with $V_{\max}$ in both the analytical model ($0.58 \rightarrow 0.47$, $r = -0.87$) and real eddies ($2.35 \rightarrow 0.35$, $r = -0.96$). The ablation experiment quantifies uncertainty costs, decomposes marginal effects, detects negative synergy, and characterizes variance reduction.
3. Methodology transfer. The ACMS algorithm and parameter ranking methodology transfer to real eddy environments. ACMS converges on all four HYCOM eddies ($\rho_S > 0.95$), with vertical profile metrics universally retained. Parameter importance rankings are statistically consistent across environments (Kendall $W = 0.59$, $p = 0.037$), with η_v consistently ranked first.

Supplementary Materials: The following supporting information can be downloaded at: <https://www.mdpi.com/article/10.3390/jmse14131221/s1>, Section S1: Fleet-Level Evaluation Metric System (Table S1: 27 fleet-level candidate metrics; Table S2: 9 core metrics selected by ACMS); Section S2: ACMS Backward Stepwise Elimination Trace (Table S3: 18-step elimination trace); Section S3: Five-Fold Cross-Validation of ACMS (Table S4: cross-validation results); Section S4: Weight Sensitivity Analysis (Tables S5–S7: weighting schemes, score correlations, ablation robustness); Section S5: Multi-Criterion Consensus Boundary Identification; Section S6: Per-Eddy ACMS Core Metric Lists (Tables S8–S12: per-eddy core metrics, cross-eddy stability, parameter sensitivity) [51–55].

Funding: This research was funded by the National Natural Science Foundation of China (Grant Nos. 42506182, 52405287 and 52505033), the Joint Funds of Natural Science Foundation of Tianjin (Grant No. 25JCLQJC00190), the Natural Science Foundation of Tianjin (Grant No. 25JCYBJC01650), the 2026 Tianjin Metrology Science and Technology Project (Grant Nos. 2026TJMT024, 2026TJMT061, 2026TJMT053), and the Tianjin Municipal Education Commission Scientific Research Program Project (Grant No. 2024KJ082).

Author Contributions: Conceptualization, D.X. and Y.Z.; methodology, D.X.; software, D.X.; validation, D.X. and S.L. (Shuai Li); formal analysis, D.X.; investigation, D.X., F.L., S.L. (Shuai Li), and S.L. (Shufeng Li); resources, Y.Z.; data curation, D.X. and X.W.; writing—original draft preparation, D.X.; writing—review and editing, Y.Z. and P.W.; visualization, D.X.; supervision, Y.Z.; project administration, Y.Z.; funding acquisition, Y.Z. and D.X. All authors have read and agreed to the published version of the manuscript.

Institutional Review Board Statement: Not applicable.

Informed Consent Statement: Not applicable.

Data Availability Statement: The HYCOM reanalysis data used in this study are publicly available at <https://www.hycom.org/dataserver/gofs-3pt1/analysis> (accessed on 2 May 2026). Simulation code and processed data are available upon request from the corresponding author.

Conflicts of Interest: The authors declare no conflicts of interest. The funders had no role in the design of the study; in the collection, analyses, or interpretation of data; in the writing of the manuscript; or in the decision to publish the results.

Abbreviations

The following abbreviations are used in this manuscript:

ACMS	Adaptive Core Metric Selection.
AUV	Autonomous Underwater Vehicle.
CV	Coefficient of Variation.
HYCOM	HYbrid Coordinate Ocean Model.
LHS	Latin Hypercube Sampling.
MI	Mutual Information.
NGC	Normalized Gap Closure.
OLS	Ordinary Least Squares.
SCOPE	Stochastic Coordinated Observation Performance Evaluation.
SSH	Sea Surface Height.

References

1. Chelton, D.B.; Schlax, M.G.; Samelson, R.M. Global observations of nonlinear mesoscale eddies. *Prog. Oceanogr.* **2011**, *91*, 167–216.
2. Dong, C.; You, Z.; Dong, J.; Ji, J.; Sun, W.; Xu, G.; Lu, X.; Xie, H.; Teng, F.; Liu, Y.; et al. Oceanic Mesoscale Eddies. *Ocean-Land Res.* **2025**, *4*, 0081.
3. Zhang, Z.; Wang, W.; Qiu, B. Oceanic mass transport by mesoscale eddies. *Science* **2014**, *345*, 322–324.
4. McGillicuddy, D.J. Mechanisms of physical-biological-biogeochemical interaction at the oceanic mesoscale. *Annu. Rev. Mar. Sci.* **2016**, *8*, 125–159.

5. Dong, C.; Liu, L.; Nencioli, F.; Bethel, B.J.; Liu, Y.; Xu, G.; Ma, J.; Ji, J.; Sun, W.; Shan, H.; et al. The near-global ocean mesoscale eddy atmospheric-oceanic-biological interaction observational dataset. *Sci. Data* **2022**, *9*, 436.
6. Bian, C.; Jing, Z.; Wang, H.; Wu, L.; Chen, Z.; Gan, B.; Yang, H. Oceanic mesoscale eddies as crucial drivers of global marine heatwaves. *Nat. Commun.* **2023**, *14*, 2970.
7. Toyoda, T.; Nakano, H.; Urakawa, L.S.; Sugiyama, T.; Aoki, K.; Nishikawa, S.; Ishikawa, Y.; Usui, N.; Nagano, A.; Osafune, S.; et al. Mesoscale eddies with North Pacific subtropical mode water as observed by an autonomous underwater vehicle during April–June 2018. *Prog. Oceanogr.* **2026**, *245*, 103743.
8. Li, S.; Cao, X.; Ma, W.; Liang, Y.; Xue, D. Taking into Account the Eddy Density on Analysis of Underwater Glider Motion. *J. Mar. Sci. Eng.* **2022**, *10*, 1638.
9. Chaigneau, A.; Le Texier, M.; Eldin, G.; Grados, C.; Pizarro, O. Vertical structure of mesoscale eddies in the eastern South Pacific Ocean. *J. Geophys. Res.* **2011**, *116*, C11025.
10. Rudnick, D.L. Ocean Research Enabled by Underwater Gliders. *Annu. Rev. Mar. Sci.* **2016**, *8*, 519–541.
11. Testor, P.; de Young, B.; Rudnick, D.L.; Glenn, S.; Hayes, D.; Lee, C.M.; Pattiaratchi, C.; Hill, K.; Heslop, E.; Turpin, V.; et al. OceanGliders: A component of the integrated GOOS. *Front. Mar. Sci.* **2019**, *6*, 422.
12. He, B.; Zhang, R.; Pan, R.; Wang, J.; Wan, X.; Sang, J. Underwater glider fleet control and application: State of the art and future prospect. *Ocean* **2026**, 9470023.
13. Chen, Y.; Li, Q.P.; Yu, J. Submesoscale variability of subsurface chlorophyll-a across eddy-driven fronts by glider observations. *Prog. Oceanogr.* **2022**, *207*, 102880.
14. Damerell, G.M.; Bosse, A.; Fer, I. Merging of a mesoscale eddy into the Lofoten Vortex captured by an ocean glider and SWOT. *Ocean Sci.* **2025**, *21*, 2763–2785.
15. Xue, D.Y.; Wu, Z.L.; Wang, Y.H.; Wang, S.X. Coordinate control, motion optimization and sea experiment of a fleet of Petrel-II gliders. *Chin. J. Mech. Eng.* **2018**, *31*, 17.
16. Zang, W.; Song, D.; Yao, P. Cooperative Sampling Path Planning of Underwater Glider Fleet. *J. Ocean Univ. China* **2023**, *22*, 907–918.
17. Li, S.; Wang, S.; Zhang, F.; Wang, Y. Glider Networking for Submesoscale Eddy Observation: An Unsupervised Approach to Sampling Point Optimization. *J. Mar. Sci. Eng.* **2025**, *13*, 1356.
18. Su, W.; E, X.; Jing, Z.; Chen, S.X. Glider path design and control for reconstructing three-dimensional structures of oceanic mesoscale eddies. *Proc. R. Soc. A* **2026**, *482*, 20250533.
19. Qin, Y.; Guan, S. Motion Simulation of Spatial Sampling of Mesoscale Processes for Underwater Gliders. *J. Unmanned Undersea Syst.* **2022**, *30*, 474–484.
20. Zhao, W.; Yu, J.; Zhang, F.; Huang, Y.; Zhang, A. Tracking moving mesoscale eddies with underwater gliders under autonomous prediction and control. *Control Eng. Pract.* **2021**, *113*, 104839.
21. Ma, X.; Wang, Y. The Adaptive Sampling of Marine Robots in Ocean Observation: An Overview. *IEEE J. Ocean. Eng.* **2025**, *50*, 1103–1126.
22. Song, D.; Gan, W.; Yao, P. Search and tracking strategy of autonomous surface underwater vehicle in oceanic eddies based on DRL. *Appl. Soft Comput.* **2023**, *132*, 109902.
23. Zhang, R.; He, B.; Wang, Y.; Ma, W.; Yang, S. Recent advances in path planning for underwater gliders: A comprehensive review. *Ocean Eng.* **2024**, *299*, 117166.
24. Merci, A.; Anthierens, C.; Thirion-Moreau, N.; Le Page, Y. A simulator of underwater glider missions for path planning. *Ocean Eng.* **2022**, *269*, 113514.
25. Wu, H.; Niu, W.; Wang, S.; Yan, S. Sensitivity analysis of control parameter errors and current parameters to motion accuracy of underwater glider using Sobol' method. *Appl. Ocean Res.* **2021**, *110*, 102626.
26. Wu, H.; Niu, W.; Wang, S.; Yan, S. A feedback control strategy for improving the motion accuracy of underwater gliders in currents. *Ocean Eng.* **2022**, *252*, 111250.
27. Lan, W.; Jin, X.; Chang, X.; Wang, T.; Zhou, H.; Tian, W.; Zhou, L. Path planning for underwater gliders in time-varying ocean current using DRL. *Ocean Eng.* **2022**, *262*, 112226.
28. Lan, W.; Jin, X.; Chang, X.; Zhou, H. Based on Deep Reinforcement Learning to path planning in uncertain ocean currents for Underwater Gliders. *Ocean Eng.* **2024**, *301*, 117501.
29. Hu, H.; Zhang, Z.; Li, L.; Peng, X. Energy-optimal motion planning of underwater gliders accounting for seabed topography and ocean currents. *Ocean Eng.* **2023**, *288*, 116008.
30. Song, Y.; Ye, H.; Wang, Y.; Niu, W.; Wan, X.; Ma, W. Energy Consumption Modeling for Underwater Gliders Considering Ocean Currents and Seawater Density Variation. *J. Mar. Sci. Eng.* **2021**, *9*, 1164.
31. Lermusiaux, P.F.J. Adaptive modeling, adaptive data assimilation and adaptive sampling. *Phys. D* **2007**, *230*, 199–217.
32. Duan, Y.; Zhang, H.; Chen, X.; Zhou, M. A Gaussian Function Model of Mesoscale Eddy Temperature Anomalies and Research of Spatial Distribution Characteristics. *Remote Sens.* **2024**, *16*, 1716.

33. Meunier, T.; Ménesguen, C.; Schopp, R.; Le Gentil, S. Tracer stirring around a meddy. *J. Mar. Res.* **2015**, *73*, 1–19.
34. Ferrari, R.; Wunsch, C. Ocean circulation kinetic energy: Reservoirs, sources, and sinks. *Annu. Rev. Fluid Mech.* **2009**, *41*, 253–282.
35. Carton, X. Hydrodynamical modeling of oceanic vortices. *Surv. Geophys.* **2001**, *22*, 179–263.
36. Leonard, N.E.; Paley, D.A.; Davis, R.E.; Fratantoni, D.M.; Lekien, F.; Zhang, F. Coordinated control of an underwater glider fleet in an adaptive ocean sampling field experiment. *J. Field Robot.* **2010**, *27*, 718–740.
37. Zamuda, A.; Hernández Sosa, J.D.; Adler, L. Constrained differential evolution optimization for underwater glider path planning in sub-mesoscale eddy sampling. *Appl. Soft Comput.* **2016**, *42*, 441–457.
38. Zang, W.; Yao, P.; Song, D. Standoff tracking control of underwater glider to moving target. *Appl. Math. Model.* **2022**, *102*, 485–502.
39. Zhao, W.; Zhou, Y.; Yu, J.; Zhang, A.; Huang, Y. Tracking control of underwater gliders in ocean mesoscale eddies observation task. In Proceedings of OCEANS 2016 MTS/IEEE Monterey, Monterey, CA, USA, 19–23 September 2016; pp. 1–5.
40. Li, S.; Zhang, F.; Wang, S.; Wang, Y.; Yang, S. Constructing the three-dimensional structure of an anticyclonic eddy with the optimal configuration of an underwater glider network. *Appl. Ocean Res.* **2020**, *95*, 101893.
41. Helton, J.C.; Davis, F.J. Latin hypercube sampling and the propagation of uncertainty. *Reliab. Eng. Syst. Saf.* **2003**, *81*, 23–69.
42. Saltelli, A.; Ratto, M.; Andres, T.; Campolongo, F.; Cariboni, J.; Gatelli, D.; Saisana, M.; Tarantola, S. *Global Sensitivity Analysis: The Primer*; Wiley: Chichester, UK, 2008.
43. Kraskov, A.; Stögbauer, H.; Grassberger, P. Estimating mutual information. *Phys. Rev. E* **2004**, *69*, 066138.
44. Cohen, J. *Statistical Power Analysis for the Behavioral Sciences*, 2nd ed.; Lawrence Erlbaum Associates: Hillsdale, NJ, USA, 1988.
45. Sawilowsky, S.S. New effect size rules of thumb. *J. Mod. Appl. Stat. Methods* **2009**, *8*, 597–599.
46. Wilcoxon, F. Individual comparisons by ranking methods. *Biom. Bull.* **1945**, *1*, 80–83.
47. Beyer, H.-G.; Sendhoff, B. Robust optimization—A comprehensive survey. *Comput. Methods Appl. Mech. Eng.* **2007**, *196*, 3190–3218.
48. Chassignet, E.P.; Hurlburt, H.E.; Smedstad, O.M.; Halliwell, G.R.; Hogan, P.J.; Wallcraft, A.J.; Baraille, R.; Bleck, R. The HYCOM (HYbrid Coordinate Ocean Model) data assimilative system. *J. Mar. Syst.* **2007**, *65*, 60–83.
49. Metzger, E.J.; Smedstad, O.M.; Thoppil, P.G.; Hurlburt, H.E.; Cummings, J.A.; Wallcraft, A.J.; Zamudio, L.; Franklin, D.S.; Posey, P.G.; Phelps, M.W.; et al. US Navy operational global ocean and Arctic ice prediction systems. *Oceanography* **2014**, *27*, 32–43.
50. Kendall, M.G.; Babington Smith, B. The problem of m rankings. *Ann. Math. Stat.* **1939**, *10*, 275–287.
51. Le Vu, B.; Stegner, A.; Arsouze, T. Angular momentum eddy detection and tracking algorithm (AMEDA) and its application to coastal eddy formation. *J. Atmos. Ocean. Technol.* **2018**, *35*, 739–762.
52. Nencioli, F.; Dong, C.; Dickey, T.; Washburn, L.; McWilliams, J.C. A vector geometry-based eddy detection algorithm and its application to a high-resolution numerical model product and high-frequency radar surface velocities in the Southern California Bight. *J. Atmos. Ocean. Technol.* **2010**, *27*, 564–579.
53. Okubo, A. Horizontal dispersion of floatable particles in the vicinity of velocity singularities such as convergences. *Deep-Sea Res.* **1970**, *17*, 445–454.
54. Rasmussen, C.E.; Williams, C.K.I. *Gaussian Processes for Machine Learning*; MIT Press: Cambridge, MA, USA, 2006.
55. Weiss, J. The dynamics of enstrophy transfer in two-dimensional hydrodynamics. *Physica D* **1991**, *48*, 273–294.

Disclaimer/Publisher’s Note: The statements, opinions and data contained in all publications are solely those of the individual author(s) and contributor(s) and not of MDPI and/or the editor(s). MDPI and/or the editor(s) disclaim responsibility for any injury to people or property resulting from any ideas, methods, instructions or products referred to in the content.