



Article The Lower Havel River Region (Brandenburg, Germany): A 230-Year-Long Historical Map Record Indicates a Decrease in Surface Water Areas and Groundwater Levels

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Abstract: Instrumental data show that the groundwater and lake levels in Northeast Germany have decreased over the past decades, and this process has accelerated over the past few years. In addition to global warming, the direct influence of humans on the local water balance is suspected to be the cause. Since the instrumental data usually go back only a few decades, little is known about the multidecadal to centennial-scale trend, which also takes long-term climate variation and the long-term influence by humans on the water balance into account. This study aims to quantitatively reconstruct the surface water areas in the Lower Havel Inner Delta and of adjacent Lake Gülpe in Brandenburg. The analysis includes the calculation of surface water areas from historical and modern maps from 1797 to 2020. The major finding is that surface water areas have decreased by approximately 30% since the pre-industrial period, with the decline being continuous. Our data show that the comprehensive measures in Lower Havel hydro-engineering correspond with groundwater lowering that started before recent global warming. Further, large-scale melioration measures with increasing water demands in the upstream wetlands beginning from the 1960s to the 1980s may have amplified the decline in downstream surface water areas.

Keywords: long-term hydrological changes; historical maps; review of written sources; preindustrial to industrial period; hydro-engineering history; effects of global warming; drying trend; wetlands; drainage works to create cropland; Lower Havel River Region; Brandenburg; Germany

1. Introduction

1.1. Droughts in Lowland River Floodplains and Peatlands

Lowland river floodplains and Central European peatlands are especially sensitive to increasing drought periods in the context of global warming [1–7]. In the already relatively dry continental regions in Germany's Federal State of Brandenburg, groundwater and lake levels have declined significantly over the past 40–50 years, in some cases by up to 3 m [8–10]. In addition to the current global warming which constitutes an indirect human effect [11], direct human impacts via hydro-engineering and land reclamation measures



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Copyright: © 2022 by the authors. Licensee MDPI, Basel, Switzerland. This article is an open access article distributed under the terms and conditions of the Creative Commons Attribution (CC BY) license (https:// creativecommons.org/licenses/by/ 4.0/). have also been discussed as potential factors [12]. In particular, the extensive drainage works and groundwater regulation measures from the 1960s to the 1980s and water-sapping pine monocultures have been mentioned in this context [13,14]. Scenarios also suggest that water availability will continue to deteriorate significantly in the future [15].

However, the empirical findings on decreasing water quantities in Brandenburg and neighboring regions are mostly based on instrumental records of river, groundwater and lake levels, which usually only cover the last 30–60 years [9]. Only in very few cases do the records go beyond this period, often concerning important rivers for inland navigation (e.g., Elbe and Oder). However, their levels can be considerably disturbed by hydro-engineering and cannot be used for regional upscaling. Except in the few recorded cases of extreme flood events [16], the "societal memory" and thus the region-specific knowledge about hydrological variability are consequently limited to the period observed by instruments.

1.2. Systematic Quantitative Analyses of Historical Maps for the Reconstruction of Hydrological Dynamics in Floodplains and Wetlands

The reconstruction of paleohydrological processes over the past centuries is increasingly relying on historical maps. This method can produce valid data [17] and extend the time frame beyond the duration of instrumental records. Among the hydrological dynamics that can be addressed are long-term lake level fluctuations [9,17–22], past river courses [23–26], surface water areas in floodplains [27], changes in floodplain geomorphologies [28,29], multi-centennial changes in aquatic ecological connectivities [30] and complex socio-hydrological models [17,31,32]. These studies focus on the exploration of the long-term consequences of hydraulic engineering interventions by humans [33,34] and of the possible influence of climatic changes on the water balance.

1.3. Aims of the Study

This study aims to quantitatively reconstruct the human impact on changes in surface water areas and groundwater levels within the Lower Havel River Region in Germany's Federal State of Brandenburg. To this end, a multi-proxy approach is used, which allows insights into paleohydrological processes from the 13th century and thus well beyond the time period of instrumental measurements, which only span the period from 1950 to 2021. On the regional scale, the study first conducts a review of written sources reporting historical human measures for the Lower Havel River Region. Here, the focus is on hydroengineering measures in floodplains on the one hand and on land reclamation, drainage works and melioration measures in peatlands on the other. In the next step, this historical information is processed with a semi-quantitative spatiotemporal approach to capture spatial hotspots and temporal peaks of human forcing in the study region. This data set forms a standardized basis for documenting human activities that have had a potential impact on changes in surface water areas in the region since the late medieval period.

On the local scale, the changes in surface water areas are then reconstructed for two smaller study sites using a 230-year-long record derived from historical and topographic maps. This data set is then compared to the historical record from the review. Our multiproxy approach enables us to detect direct anthropogenic forcing mechanisms affecting potential changes in surface water areas during the last two centuries. In a final step, we compare our findings to long-term temperature records to discuss the potential impact of climatic parameters on changes in surface water areas.

2. Materials and Methods

2.1. The Region-Scale Study Area: The Lower Havel River Region

The Havel River has a catchment area of approximately 24,000 km² and a length of 325 km. It is a lowland river with a small difference in altitude between its source (63 m a.s.l.) and mouth (22 m a.s.l.) at the Elbe River. The Havel River flows through numerous lakes and is partially characterized by a multi-channel system [35]. The Lower

Havel River Region (Figure 1) represents the western part of the Brandenburg glacial area [36], which consists of gently undulating till plains, sandy valley fills with covering dunes and hilly terminal moraines [37]. The region is characterized by a large proportion of fens. Of particular importance here are the Rhinluch and the Havelländisches Luch ('Luch' means swamp or fen) (Figure 1). The large fens lie in former glacial valleys, which were no longer active at the end of the Weichselian glacial period. Afterward, isolated kettle lakes formed, some of which had already silted up by the late-glacial to early Holocene. In the Holocene, large areas of peat began to form on the sandy deposits of the glacial valleys, which ended with fen formation in the late Holocene [38–40].

Human settlement in the Lower Havel River Region began in the Final Paleolithic at the latest. The human activities in the subsequent Mesolithic did not result in any lasting changes to the landscape [41-45]. Rather, what changed the landscape were the sea level rises during the Calais transgressions in the Middle Holocene [46,47], which led to the formation of backwater of the Havel and, in turn, to an influx of Elbe water into the Havel lowlands and ultimately to excessive fen formation [38,48]. With the transition to the Neolithic [49] and the subsequent Bronze and Iron Ages, there was initial deforestation and evidence of the first cultivated plants, although the grassland open habitats that are still characteristic today in the Havel floodplain are associated with a natural origin in the early Holocene [50,51]. Slavic population groups began to settle in the area from the 7th century AD onward, but the high groundwater level and frequent flooding in the area of the Havel floodplain and adjacent Rhinluch fen beginning in the Subboreal [40,43] threatened settlements [52]. Slavic fortifications and settlements were concentrated along watercourses and lakes and were therefore particularly exposed to hydrological changes [53,54]. Since the beginning of the high medieval Ostsiedlung (German colonization, c. 1150–1300 AD), the Elbe River and the Lower Havel River Region have seen a radical phase of landscape and cultural change [55]. This includes large-scale clearing, the foundation of villages and towns, the construction of water mill dams at the towns of Rathenow and Brandenburg and the resulting changes in the water balance [56–58]. Another important fact concerning the lower Havel River is that all of the shipping traffic during the Middle Ages from the Mark Brandenburg toward the Port of Hamburg passed through this area. The export of grain over the Havel to Hamburg was a major component of the local economy and important for the development of Berlin. In the medieval and modern period, the lower Havel was one of the main routes of traffic in the region [57,59–61]. Along the lower Havel River, the largest river and floodplain renaturation project in Central Europe is currently taking place. In an area of around 19,000 ha that extends for 90 km, the river has been given a largely "natural" hydrological regime again. This includes the restoration of floodplains, the establishment of alluvial forests, the removal of bank constructions, the reconnection of oxbow lakes, and adjustments to land use and water management [62,63].

2.2. Semi-Quantitative Approach for the Evaluation of Past Hydro-Engineering Measures and Land Reclamation Activities in the Lower Havel River Region

As a first step, selected written sources about past hydro-engineering measures in floodplains and about land reclamation activities in wetlands in the Lower Havel River Region were reviewed. The evidence is documented in Tables 2–5 and includes hydro-engineering measures for inland navigation and hydropower installations (Tables 2–4) as well as land reclamation and drainage work in wetlands (Table 5). In the second step, the human impact in the Lower Havel River Region is evaluated using a semiquantitative approach. The spatial positions of the punctual, linear and areal hydro-engineering activities are allocated to individual grids of $6 \times 6 \text{ km}^2$ (Figure 2). This approach results in a spatial and temporal distribution of the individual measures at a grid-scale resolution. Locations from written sources (Tables 2–5) are assigned to 208 squares measuring $6 \times 6 \text{ km}^2$. Every mention in a written source is considered a hydro-engineering event and recorded with the value n = 1 per associated square. Spatially larger measures are distributed over several squares. In this case, values of n = 1 are taken into account for



each affected square. Finally, all recorded hydro-engineering events and land reclamation measures are summed up per square.

Figure 1. Study area of the Lower Havel River Region. The blue arrows indicate flow directions, and the yellow dot indicates the position of the Wusterhausen (Dosse) groundwater-level measuring point. The white dots indicate hydro-engineering measures from the 13th to the 21st century (Tables 2–4). The white dots with a "W" indicate land reclamation, drainage works and melioration measures in peatlands and floodplains since the 17th century (Table 5). The yellow rectangles indicate the "Lower Havel Inner Delta" and "Lake Gülpe area" local study sites. Data source of the digital elevation model: [64] and references therein.



Figure 2. Structural concept of the semi-quantitative approach used in this study to evaluate evidence of hydro-engineering measures and land reclamation activities from the 13th to 20th centuries. Data source of the digital elevation model: [64] and references therein.

2.3. Local Study Sites for Surface Water Area Reconstructions from Maps

The local study sites the "Lower Havel Inner Delta" and the adjacent "Lake Gülpe" (yellow rectangles in Figure 1) lie in the area of the Lower Havel River Region. The large expanse of the Lower Havel Inner Delta results from the former confluence and the associated large-scale glacio-fluvial erosion processes of the Weichselian Baruth, Berlin and Eberswalde glacial valleys [36,65]. The Lower Havel Inner Delta is characterized by a multi-channel and meandering river system with oxbow lakes and episodic floods, especially during spring [66]. Lake Gülpe, which adjoins the Lower Havel Inner Delta to the east, probably owes its hollow shape to the postglacial dead ice thawing [38]. The lake is traversed by the Rhin River and reaches a maximum depth of approximately 3 m near the outlet [66]. The local study sites are part of one of a few remaining relatively "near-natural" floodplain areas in Central Europe [67] with high biodiversity and the

presence of rare plant and animal species [68,69]. Endangered migratory fish species such as asp (*Aspius aspius*) and zope (*Abramis ballerus*) spawn in the area [66]. The lake is a bird sanctuary of supra-regional importance and was designated as a nature reserve in 1967 [70]. The Lower Havel Inner Delta and the Lake Gülpe area are home to an impressive number of rare and threatened plant species and vegetation types. The latter include floodplain meadows of the alliances *Cnidion dubii* and *Potentillion anserinae* as well as herbaceous pioneer vegetation [50,71,72], which in turn contain populations of river corridor plant species that reach the western border of their European distribution area in this region [75].

2.4. Exploration of Historical and Topographic Maps for the Reconstruction of Surface Water Area Changes during the Last 230 Years

The State Survey and Geobasis Information Brandenburg [76] provides an overview of the easily accessible historic, topographic and digital maps of the Lower Havel River Region. The period from which historical sources are available spans almost the last six centuries (Table 1). An important criterion for the spatiotemporal analysis of historical maps is the possibility of georeferencing as precisely as possible. The map series noted in Table 1 with an asterisk meet this requirement. Some map series obtained as a Web Map Service (WMS) were already georeferenced. In the case of the Prussian original map series, georeferencing could be carried out with sufficient accuracy due to suitable benchmarks (e.g., churches and bridges) and coordinate grids. With the help of Quantum Geographical Information System (QGIS) version 3.22.1, the surface water areas of the map sheets from 1787, 1843, 1880, 1997 and 2020 were digitized and quantified. The oldest map series digitized was the Schmettau map series, which was created under the direction of the Prussian officer and cartographer Friedrich Wilhelm Karl Graf von Schmettau in the years 1767–1787 for the Prussian territory east of the Weser River and in the years 1780–1793 for Mecklenburg. In the context of georeferencing via topographical benchmarks and a subsequent quality check, it was possible to show that the Schmettau maps are surprisingly accurate for their time of origin [77]. For the Prussian territory, 270 sheets with a scale of 1:50,000 were created, which are now in the holdings of the Berlin State Library [78]. The original map sheet from 1843 is available as a georeferenced layer [76,79,80].

The Lower Havel Inner Delta and the Lake Gülpe area (Figure 1) were selected as the local study areas. While the first area represents a river landscape, a stagnant open water body marks the second area. Due to the specific behaviors of each hydro-geomorphic system, the quality of the map sheet series and the methodological approach were checked by analyzing the local areas separately and then comparing the results with one another.

In order to derive lake-level fluctuations from the changing surface water areas, transects (T1–T5) were created at the margins of Lake Gülpe, in which the lake shorelines read from the maps were linked to the heights from the Light Detection and Ranging (LiDAR) digital terrain model [76]. Only the shorelines of the maps from 1843 AD, 1880 AD, 1997 AD and 2020 AD [76] were taken into account, since the Schmettau map series from 1787 AD [76] was found to not have sufficient positional accuracy for the conversion into lake levels.

Table 1. Overview of historical and modern maps from the Lower Havel River Region. Maps were marked with an asterisk (*) if they were found to be suitable for our reconstruction of past landscape dynamics according to position accuracy. This was also possible with some restrictions for the Schmettau map series. WMS: Web Map Service.

Map Series	Year	Scale
Brandenburgensis Marchae Descriptio [76]	1588	1:880,000
Markgrafschaft und Kurfürstentum Brandenburg [76]	1696	1:450,000
Karte des Kurfürstentums Brandenburg [76]	1750	1:550,000

Map Series	Year	Scale
Geographische Karte des Kurfürstentums Brandenburg [76]	1758	1:550,000
Neue geographische Karte der Markschaft Brandenburg [76]	1773	1:625,000
Schmettausches Kartenwerk Brandenburg (Schmettau map series) (*): WMS [76]	1787	1:50,000
Spezialkarte von der Mittelmark [76]	1790	1:300,000
Preußische Urmesstischblätter (Prussian original map series) * [76]	1843	1:25,000
Geologische Karte (GK25) (Geological map series) * [81]	1880 (1883–1884)	1:25,000
Deutsches Reich-Messtischblätter (Topographic map series) *: WMS [76]	1880 (1882)	1:25,000
Digital topographic map (preliminary edition) *: WMS [76]	1997	1:10,000
Digital topographic map *: WMS [76]	2020	1:10,000
Light Detection and Ranging (LiDAR)-based digital terrain model* [76]	2020	1-m resolution

Table 1. Cont.

3. Results

3.1. Review of the History of Human Impacts on the Region-Scale Water Balance

Written sources contained ample information on the settlement history and water regulation measures in the study region. We found that, in the course of the medieval Ostsiedlung (German colonization, c. 1150–1300 AD), two water mills were built in the cities of Rathenow and Brandenburg at the lower Havel River (Table 2). Since the 13th century AD, the weirs of water mills in combination with fish weirs apparently caused a general increase in river levels of 1–2 m along the Havel River and in adjacent rivers and streams [56,60,82,83]. These processes also led to further expansion of the fens near the Havel riverbank. These same weirs also significantly restricted inland navigation [82]. In the 13th century, the Elbe dike pushed back the floods of the Elbe from the Havel floodplain [38], but at the same time, there was frequent pondage of the Havel River itself, which led to further waterlogging [37]. With regard to the cities' need for grain, the rights to damming mills had priority over the interests of agriculture for many centuries [59]. Thus, human activities up through the early modern period intensified the natural character of the landscape as a boggy lowland area.

Table 2. Hydro-engineering projects in the Lower Havel River Region from the 13th to 17th centuries AD. Numbers refer to locations on Figure 1.

No	Construction Activities (Hydro-Engineering)	Construction Time
1	Brandenburg: mill dam [56,58,60]	Appox. 1200
2	Brandenburg: shipping pier [57]	1274
3	Rathenow: mill weir [61]	1288
4	Brandenburg: mill weir [58,61]	1309
5	Rathenow: moat [57]	1288
6	Brandenburg: flood channel for a mill [57]	1315
7	Brandenburg: water mill [84]	1326
8	Rathenow: mill weir [84]	1335
9	Havelberg: flood channel for a mill [57]	1375
10	Havelberg: two floating mills [57]	1375
11	Lower Havel and tributaries: multiple fish weirs [84]	1375
12	Brandenburg: drawbridge for ship passages [57]	1423
13	Brandenburg: bank reinforcement [84]	1455

No	Construction Activities (Hydro-Engineering)	Construction Time
14	Brandenburg: navigation lock [57]	1548–1550
15	Rathenow: navigation lock [57]	1548–1559
16	Rathenow: flood channel system for mills [57]	Approximately 1550

Table 2. Cont.

The abundance of water, however, was countered with great efforts in the following centuries [8,37]. First, land reclamation measures were undertaken in the 17th century AD [59,85]. Specialists in the construction of dikes, the creation of drainage ditches and the market-oriented introduction of dairy farms came from Holland [59]. In 1718, Prussian King Friedrich Wilhelm I began the reclamation of large fen areas of the Rhinluch and the Havelländisches Luch [86] (Table 5) and the settlement of people ("Peuplierung") there. The construction of the Havelländische Große Hauptkanal took place between 1718 and 1725 [87] (Table 3 and Figure 1), during which the local population also had to provide auxiliary services [59]. The Königshorst domain became an agricultural model farm that supplied the royal court.

Table 3. Hydro-engineering projects in the Lower Havel River Region during the 18th and 19th centuries AD. Numbers refer to locations on Figure 1.

No	Construction Activities (Hydro-Engineering)	Construction Time
17	Construction of the Havelländischer Großer Hauptkanal [59]	1718–1725
18	Potsdam: city canal [82]	1770
19	Plaue, Pritzerbe, Bahnitz: removal of weirs to improve inland navigation [88]	1771
20	Sandau: Elbe-Havel embankment [61]	1771–1772
21	Sandau: strengthening of the Elbe-Havel embankment [82]	1809
22	Sandau: strengthening of the Elbe-Havel embankment [59]	1832
23	Quitzöbel: strengthening of the Elbe-Havel embankment [82]	1832–1836
24	Quitzöbel: prolongation of the Elbe-Havel embankment [61]	1832–1836
25	Quitzöbel: relocation of the Havel mouth [61]	1832–1836
26	Between Brandenburg and Rathenow: removal of 106 fish weirs to improve inland navigation [89]	1837–1842
27	Construction of the Sacrow–Paretz Canal [82,86]	1874–1878
28	Lower Havel: river straightening [82]	1875–1881
29	Lower Havel: river straightening [61,82]	1882–1898
30	Brandenburg: new navigation lock [90]	1881–1883
31	Broadening of the Sacrow-Paretz Canal [61]	1888–1890
32	Lower Havel: river straightening between Havelberg and Rathenow [61]	1897–1902
33	Rathenow: second river course for inland navigation [61]	1898–1901

Under Frederick the Great (1712–1786) in particular, further groundwater lowering measures were carried out in the Dossebruch (Dosse swamp), such as the construction of drainage ditches and the demolition of water mills on the Rhin River in 1773. In the years 1773–1775, the Friedrichsdorf colony and six other villages were built southwest of Neustadt (Dosse) [37,59] (Table 5). Settlers from the Palatinate and Rhineland were hired to live here [91]. With the drainage of the Rhinluch and Dossebruch, the landscape was intensively reshaped [37]. The conditions in the meliorated wetlands were favorable for cattle farming and the cultivation of rye, potatoes and vegetables. All the new agricultural

areas were settled with recruited people. In addition, Poles, Saxons and Mecklenburgers came into the country [91]. In total, around 300,000 colonists had immigrated to Prussia by the time of the death of Frederick the Great in 1786. Around 20,000 colonists were settled in the former Rhin, Dosse and Jäglitz wetlands [91]. The extensive land reclamation of the 18th century AD, however, had only little influence on the still-recurring and widespread flooding in the former wetlands [37].

However, in the upper Rhinluch and the Havelländisches Luch, peat cutting and the lowering of the groundwater level as part of large-scale drainage works to create cropland from the 18th century AD on led to a significant decline in the surface water area of Lake Wublitzsee and Lake Kremmen [92,93]. In 1841, Lake Kremmen had an area of 3.9 km², but by 1985, it had been reduced to two smaller residual lakes with a total surface water area of 0.8 km² [93].

Since the mid-1840s, there has been an increase in aquatic plants in the nutrient-rich waters of the Havel, which has had a negative impact on the drainage. Among these plants was the invasive Canadian waterweed (*Elodea canadensis*), which became a significant problem [59]. Limnological studies also show trends toward increasing eutrophication of Brandenburg and Mecklenburg water bodies beginning in the first half of the 20th century at the latest [94].

The natural flood frequency within the lower Havel and Dosse floodplains was altered to a great extent in the 19th and 20th centuries AD through hydraulic engineering (Tables 3 and 4). These measures have resulted in the reduction of the area of the land prone to flooding by 90% over the second half of the 20th century [37]. As a result of the hydraulic engineering work on the lower Havel (Gvensdorfer Vorfluter, embankments), the danger of Elbe floods and the Havel backwater decreased significantly from 1950 to 1960 [59,61,95]. Furthermore, the amplitude between the mean low waters and mean high waters was significantly reduced at Brandenburg (Havel) over the course of the second half of the 20th century AD [56].

Table 4. Hydro-engineering projects in the in the Lower Havel River Region during the 20th century AD. Numbers refer to locations shown in Figure 1.

No	Construction Activities (Hydro-Engineering)	Construction Time
34	Weißer See-Jungfernsee: broadening of the canal [82]	1903–1905
35	Brandenburg: enlargement of the navigation lock [90]	1904–1906
36	Lower Havel: broadening of the river course between Molkenberg and Garz [82]	1907–1909
37	Brandenburg: construction of the Silo canal [82]	1907
38	Warnau: cut of the Havel River course [82]	1908
39	Molkenberg: offset trench [82]	1908
40	Brandenburg: river embankment [90]	1907–1910
41	Brandenburg: construction of a moat [90]	1909–1910
42	Lower Havel between Havelberg and Pritzerbe:channel cuts and flood channels for inland navigation and flood management [82]	1912
43	Barrages and navigation locks at Bahnitz, Grütz and Garz [82]	1911–1912
44	Warnau: river straightening [61]	1912
45	Rathenow and Grütz: deepening of the channel bed [61]	1914–1915
46	Dosse and Jäglitz Rivers: stream engineering [35,37]	1929–1933
47	Relocation of the Rhin mouth [35,37]	1930–1934
48	Enlargement of the Havel riverbed between the Havel mouth and Garz to enhance the flow velocity of Havel summer floods [96]	1933–1935

No	Construction Activities (Hydro-Engineering)	Construction Time
49	Lower Havel: relocation of the mouth of the Havel (Gnevsdorfer Vorfluter), prevention of Elbe floods and Havel backwater [59,95]	1935–1956
50	Lower Havel: new navigation lock at Havelberg [61,96]	1933–1936
51	Lower Havel: weirs for navigation at Quitzöbel [61,96]	1936–1937
52	Construction of the Havel Canal	1951–1952
53	Lower Havel: weir for navigation at Neuwerben [61,96]	1949–1954
54	Lower Havel: weir and navigation lock at Gnevsdorf, prevention of Elbe floods and Havel backwater [59,61,96]	1952–1954
55	Strengthening of the Lower Havel embankment [61]	1960
56	Roskow and Weseram: broadening of meander cuts [61,96]	1965–1966
57	Ketzin: river straightening [61]	Approximately 196
58	Brandenburg: broadening of the Silo canal [96]	1970–1973
59	Brandenburg: rebuilding of the barrage with navigation lock [96]	1963–1970
60	Enlargement of the Havel riverbed between Lake Jungfernsee and Lake Plauer See [96]	Approx. 1992–1994
61	Brandenburg: rebuilding of the navigation lock [96]	1996
62	Brandenburg: enlagement of the Silo canal [96]	2003–2004
63	Lower Havel: overhaul of Neuwerben weir [96]	1998–1999
64	Lower Havel: overhaul of Gnevsdorf weir [96]	2001-2003
65	Lower Havel: overhaul of Havelberg navigation lock [96]	2001-2003
66	Lower Havel: overhaul of Grütz weir [96]	2001-2003
67	Lower Havel: overhaul of Garz weir [96]	2004
68	Rathenow: rebuilding of the mill wier [96]	2006
69	Lower Havel: rebuilding of the Bahnitz wier [96]	2006
70	Lower Havel: closure of the river course between Havelberg and Rathenow to commercial shipping, start of the Lower Havel rewilding program [63,96,97]	2008-present

Table 4. Cont.

Particularly from the 1960s until the political turn in 1989, so-called *Komplexmeliorationen* (hydromeliorations) were carried out in wetlands with great technical effort, with the aim of intensive agricultural use [8,61,98–100]. In the upper Rhinluch, these measures led to a significant decrease in the number of drainage trenches. The remaining trenches were dammed in the summer, which overall led to a more balanced groundwater level in the upper Rhinluch. However, it became necessary to supply the area with up to 20 million m³ of additional water from the Mürnitz and the upper Rhin catchment area annually [101].

Table 5. Land reclamation, drainage works and melioration measures (W = wetland) in the Lower Havel River Region since the 17th century AD. Numbers refer to locations shown in Figure 1.

No	Land Reclamation, Drainage Works and Melioration Activities	Construction Time
W1	Neuholland: foundation of Dutch settlement [59]	1659
W2	Dossebruch and Neustädtisches Luch: drainage works [59,86]	1670–1680
W3	Rhinluch: drainage measures at Linum (200 ha) [59]	1705–1710
W4	Havelländisches Luch: drainage works and foundation of Königshorst, Nordhoff, Lobeofsund, Kienberg, Hertefeld, Kuhhorst, Mangelshorst, Deutschhof, Hertefeld, Sandhorst, Ribbeckshorst, Dreibrück and Rohlandshorst [59,86]	1718–1724

No	Land Reclamation, Drainage Works and Melioration Activities	Construction Time
W5	Havelländisches Luch: peat cutting at Königshorst [59]	1746–1750
W6	Dossebruch (Dreetz): drainage works and foundation of Bartschendorf, Baselitz, Blumenaue, Fischershof, Giesenhorst, Michaelisbruch, Siegrothsbruch, Webersplan, Wilhelminenaue, Wolfsplan and Ziethensau [37,59]	1747–1778
W7	Dossebruch (Neustadt): drainage works and foundation of Friedrichsdorf and six other colonist villages [37,59]	1773–1775
W8	Dossebruch (Dreetz): demolition of the Dreetz water mill	1773
W9	Rhinluch: drainage measures [59]	1772–1776
W10	Upper Rhinluch: peat cutting at Fehrbellin and Linum [59]	1785 until the end of the 19th century
W11	Upper Rhinluch: construction of the Schwarzer Graben for peat cutting [59]	1797–1798
W12	Lower Havel floodplain: drainage works [59]	1777–1783
W13	Dosseluch, Rhinluch and Havelländisches Luch: foundation of the Luchgraben-Schauverband and cleaning of the existing drainage systems [59]	1842
W14	Golmer Bruch: draining of the fen [59]	1855
W15	Upper Rhinluch: melioration of former peat cutting areas between Kremmen and Fehrbellin, foundation of Zietenhorst and Linumhorst [59]	1912–1931
W16	Lower Havel floodplain: land reclamation by enclosure in the lower Havel floodplain (<i>Komplexmelioration</i>), changes from pastures to arable land [59,61]	1960–1970
W17	Upper Rhinluch: melioration (<i>Komplexmelioration</i>) with decrease in fen area (9%) but increase in summer groundwater levels [99,100]	1974–1976
W18	Dosse River floodplain: melioration and drainage works (<i>Komplexmelioration</i>) for the intensification of agrarian land use [61]	1968–1987

The intensification of agricultural activities in the 20th century represents a major human intervention in the region. Due to the immense nutrient input from agriculture, Lake Gülpe has changed from eutrophic to polytrophic [66]. With the partial elimination of intensive agrarian land use since the political turn of 1989, extensive nature conservation efforts have come into focus in the region [37]. An outstanding example is the riparian strip project "Lower Havelniederung between Pritzerbe and Gnevsdorf", a section of which comprises the Lower Havel Inner Delta and Lake Gülpe. The measures have been implemented since 2016 [101]. The aim of the rewilding measures is to allow the Lower Havel Inner Delta to flow through again all year round, to improve the multi-channel system through increased stream velocity and to create more habitats for rare flora and fauna in the form of floodplain forest recreation [97].

3.2. Semi-Quantitaive Approach: Region-Scale Distribution of Hydro-Engineering Activities and Drainage Works

If one considers the spatial distribution of hydro-engineering activities, hydromelioration measures and drainage works (Figure 3), a focus becomes apparent on the early urban settlement centers (Brandenburg, Rathenow), the mouth of the Havel River and on the hydromelioration and drainage work areas in Dosseluch and in particular the upper Rhinluch. In the upper Rhinluch, peat-cutting activities were also important. Basically, it can be seen that the entire area of the Lower Havel River Region has been affected by intensive hydraulic engineering activity over the past few centuries.

Table 5. Cont.



Figure 3. Semi-quantitative spatial representation of hydro-engineering measures from the 13th to 21st centuries in the Lower Havel River Region. The square values indicate the number of pieces of evidence from the reviewed written sources (see Tables 2–5). Data source of the digital elevation model: [64] and references therein.

3.3. Local-Scale Study Sites: Changes in Surface Water Areas during the Last 230 Years

We digitized the surface water areas from five different historic and modern maps for a section of the Lower Havel Inner Delta as well as for Lake Gülpe and then calculated their areas (Figures 4 and 5). From this, it becomes clear to what extent there has been a change in the position and size of the surface water areas in the study area since 1787 AD. Both sub-areas examined showed a declining trend that started in the pre-industrial period. Between 1787 and 1880 AD, the analyzed fluctuations in the size of the surface water areas were between 7.53 and 6.86 km² in the area of Lake Gülpe (Figure 6e) and between 3.52 and 3.1 km² in the Lower Havel Inner Delta (Figure 6f). There was a noticeable decline in the size of surface water areas between 1880 AD and 1997 AD from 6.86 to 4.77 km² at Lake Gülpe and from 3.1 to 2.41 km² in the Lower Havel Inner Delta. The values from 1997 AD remained roughly constant through 2020 AD. Accordingly, it can be said that the surface water areas have decreased considerably since 1787 AD.



Figure 4. Changes in Lower Havel Inner Delta surface water areas as a reconstruction from historical [76] and modern maps [76].



Figure 5. Changes in Lake Gülpe surface water areas as a reconstruction from historical [76] and modern maps [76]. The transects T1–T5 indicate the coupling of shorelines from map records [76] (1843 AD, 1880 AD, 1997 AD and 2020 AD) with LiDAR-based DEM data [76] for the reconstruction of lake level changes.

4. Discussion

4.1. The Historical Map Record Indicates a Decrease in the Surface Water Area

According to the historical map record, the size of the surface water areas decreased by 29.6% within the Lower Havel Inner Delta (Figure 6e) for the entire period from 1787 AD to 2020 AD. In the area of Lake Gülpe, the surface water area decrease was even higher—33.6%—in the same period (Figure 6f). In the two local study areas of the Lower Havel River Region, a decline in surface water area was particularly noticeable for the period from 1880 AD to 1997 AD. The roughly comparable declines for the Lower Havel Inner Delta and Lake Gülpe are striking. They might be attributed on the one hand to the hydrological coupling of both areas and on the other to the reliability of the map series, although we do not have any information about seasonality effects in the course of the topographical surveys. Other historical-geographical studies have pointed out the problem of map generalizations with regard to the determined changes in area sizes [102]. In the present study, however, this was unlikely, since the representation of Lake Gülpe in particular did not require a generalization, and the selected scale was at least 1:50,000 (Table 1) and therefore completely adequate.

The calculated lake level data (Figure 6d) also showed a continuous decline over the period from 1843 AD to 2020 AD. However, the calculated data for 1843 AD should be used with great caution, as the box plot showed a very large scatter. Considerably more reliable data are available for the three more recent calculations. This resulted in a decline in the lake levels of about 20–40 cm from 1880 AD to 1997 AD. These values match with the lake level information on the 1880 AD [76] and 1997 AD [76] maps, showing a decline of exactly 40 cm for this period. However, from here onward, this study relates exclusively to the calculation of the surface water areas, since this is the only way to compare the two

sub-study areas "Lower Havel Inner Delta" and "Lake Gülpe", and these data sets also appear to be robust for the entire period from 1787 AD to 2020 AD.

The data of the surface water area generated from the analysis of the historical maps are particularly valuable, because they significantly extend the length of instrumental data recordings into older times and thus provide potential clues as to which climatic and anthropogenic control variables may be responsible for the long-term water balance in the region. Furthermore, information about long-term changes in the sizes of water surface areas might be relevant for river rewilding programs [103].



Figure 6. Climatic and hydrological parameters: (**a**) instrumental record of European mean summer temperature anomalies relative to the 1961–1990 baseline [104], (**b**) change in global surface temperature as a decadal average [11], (**c**) instrumental record of the Wusterhausen (Dosse, Figure 1; lat: 52.883268, long: 12.46536) groundwater level [105], (**d**) reconstructed lake levels (box plots) from past Lake Gülpe shorelines from historical [76] and modern maps [76] (1843 AD, 1880 AD, 1997 AD and 2020 AD) and coupled LiDAR DEM data (own data analysis on the basis of transects T1–T5 in Figure 5), (**e**) changes in surface water area at Lake Gülpe as a reconstruction from historical [76] and modern maps [76] (1787 AD, 1843 AD, 1880 AD, 1997 AD, 2020 AD and this study) and (**f**) changes in surface water area in the Lower Havel Inner Delta as a reconstruction from historical [76] and modern maps [76] (1787 AD, 1843 AD, 1880 AD, 1997 AD, 2020 AD and own data analysis).

4.2. Climatic Effects on the Decline in Surface Water Areas

The low values of the surface water areas in the last few decades correspond to the global temperature record, which has documented higher mean temperatures since at least the 1980s (Figure 6b [11]). This observation is also supported by meteorological data on the regional and local scales. The instrumentally measured mean summer temperatures in Europe have noticeably increased in the same period (Figure 6a [104]), and the mean annual temperatures recorded by the meteorological station in Berlin-Dahlem also show an increase of around 1.0 K since the 1970s [106]. Furthermore, an available instrumental record from a groundwater-level-monitoring well at Wusterhausen (Dosse) (Figure 6c), which is 20 km away from Lake Gülpe, strongly supports the multi-decadal drying phase evidenced by the map data set. Lake level records from the Havel headwaters have exhibited decreasing levels during recent decades as well [107].

A paleohydrological study from the Schorfheide region, which is about 100 km northeast of Lake Gülpe, concluded that lake levels have fallen sharply in the last four decades [9]. The study was based on the comparison of instrumental data and historical sources such as aerial photographs and maps. The authors concluded that the lake level fluctuations in the Schorfheide area have mainly been related to the climate over the past 100 years. However, the authors pointed out that this does not apply to all lakes in the Schorfheide area and that other hydrological factors must be considered at the local level, which can also be tied to human impact (e.g., anthropogenic forest composition and drainage measures). In this context, the present study has to consider the possible human impact on the water balance in the Lower Havel River Region during the last few centuries.

4.3. Semi-Quantitaive Approach: Potential Human Impact on Surface Water Areas

The evaluation of the written sources using our semi-quantitative approach showed two phases of intense activity for the land reclamation, drainage works and melioration measures, with the first in the 18th century AD and the second in the second half of the 20th century AD (Figure 7a). The first phase was characterized by the primary land reclamation with draining activities and the foundation of colonist villages. In the second phase in the 20th century, the focus was on embankment projects as well as weir constructions for groundwater regulation. In the second phase, there was an additional need for water, as it was desired to raise the groundwater level in the fens during the summer months [59,99]. The land reclamation activities of the first phase led to a sharp decline in groundwater levels within the directly affected peatlands [19,92,93,108] but probably had a low impact on the water received by the Lower Havel Inner Delta and Lake Gülpe. Based on the analysis of the historical maps (Figure 7d,e), only a slight decline in the surface water area could be inferred. The situation was different for the melioration measures in the second phase. Both the increased water demand in the peatlands due to the damming of the drainage trenches in summer [99] and the dike construction measures in the lower Havel floodplain [61] might have been associated with the noticeable reduction in the surface water area of the Lower Havel Inner Delta and the Lake Gülpe area (Figure 7d,e).

Enhanced human hydro-engineering measures (Figure 7b) and land reclamation activities (Figure 7a) indicate an increasing human impact on the Lower Havel River Region since the 18th century AD. At a more regional scale, the pollen record from Lake Tiefer See (Klocksin lake chain in northeastern Germany) showed a noticeable increase in land use taxa since the onset of the 19th century [109,110] (Figure 7c). The lake is located in a roughly comparable region approximately 150 km away.

4.4. Potential Impacts of the Elbe Channel Incision

In addition to climatic factors and the human impact, hydrological processes must be considered as potentially having contributed to the decline in the surface water areas in the two local study sites of the Lower Havel River Region in recent decades. In the Lower Havel Inner Delta, the decline in the surface water area does not necessarily have to correlate with the water balance, as the runoff also depends on the flow speed and the depth of the channels. In this context, the current deepening of the Elbe riverbed may be an important factor. Gabriel et al. [111] documented that the incision of the Elbe riverbed is a significant process, especially between Mühlberg and the Saale mouth over a distance of c. 180 km. During the last 130 years, the riverbed has been deepened by up to 130 cm, with considerable consequences for the water balance in the Elbe floodplain. However, further downstream in the area of the Havel mouth near Havelberg, the Elbe River is not significantly affected by riverbed deepening [111]. Therefore, Elbe riverbed deepening currently has little hydrological impact on the lower Havel floodplain.



Figure 7. (a) Land reclamation, drainage works and wetland meliorations at the Lower Havel River Region described in written sources (in this study, number of evidence (*n*) per 6*6 km²), (b) hydroengineering measures in the Lower Havel River Region described in written sources (in this study, number of pieces of evidence (*n*) per 6*6 km²), (c) land use taxa (pollen counts) from the Lake Tiefer See pollen record (Figure 1a, Klocksin lake chain in northeastern Germany [109,110]), (d) changes in surface water area at Lake Gülpe as a reconstruction from historical [76] and modern maps [76] (1787 AD, 1843 AD, 1880 AD, 1997 AD, 2020 AD and own data analysis) and (e) changes in surface water area in the Lower Havel Inner Delta as a reconstruction from historical [76] and modern maps [76] (1787 AD, 1843 AD, 1880 AD, 1997 AD, 2020 AD and own data analysis).

5. Conclusions

We used a multi-proxy approach to quantitatively reconstruct long-term hydrological changes in a hotspot of human impact in Eastern Germany. Our study shows that the multi-temporal reconstruction of the sizes of surface water areas from historical and modern maps is a reliable method of quantitative analysis. Here, quantitative analyses of the surface water areas of the Lower Havel Inner Delta and the adjacent Lake Gülpe area in Brandenburg in northeast Germany were conducted. The GIS analyses for both sites were carried out independently of each other. The Lower Havel Inner Delta and the Lake Gülpe area have seen similar declines in surface water areas during the last two centuries,

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though the former is a multi-channel river and the latter is a lake. For recent decades, the generated data from the map analyses were compared with the instrumentally measured groundwater levels, and both data sets have shown a desiccation phase over the period under study.

In addition, this study presents a review of written sources that documents the spatial history of hydraulic engineering and land reclamation measures carried out across the Lower Havel River Region over the past 700 years. A grid-related, semi-quantitative approach traced the human impact on the floodplains and wetlands within the Lower Havel River Region. The Late Medieval mill dam constructions in the urban areas of Rathenow and Brandenburg led to large-scale waterlogging at the lower Havel River. With the onset of large-scale drainage works to create cropland, the hydrological conditions changed in the modern period. The semi-quantitative approach revealed that there was an intense human impact during the land reclamation, drainage works and wetland melioration phases in the 18th and 20th centuries AD and during the Lower Havel hydro-engineering measures of the second half of the 19th and the entirety of the 20th century AD. A temporal comparison of the drainage works in the wetlands and the massive hydro-engineering activities in the floodplains with the achieved surface water area reconstructions showed that the surface water areas of the two local study sites were not significantly affected by human activities between 1787 and 1880 AD. However, the surface water areas on the topographic maps from 1997 and 2020 did show a decrease by approximately 30% compared with the areas on the maps from the 18th and 19th centuries AD. This decline can be explained by the additional water demand in the peatlands due to the hydromeliorations, embankment measures in the lower Havel floodplain and the hydraulic engineering-related change in the drainage conditions in the mouth of the Havel in the second half of the 20th century. The influence of the increasingly drier climate over the past few decades must be considered as well.

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